

Is human being-soil relationship limited only to agriculture?



WHY IS SOIL IMPORTANT FOR US?

Soil functions

Soils deliver ecosystem services that enable life on Earth



2015
International
Year of Soils
fao.org/soils-2015



Food and Agriculture
Organization of the
United Nations

with the support of

Swisscontact Engagement
Centre for Swiss
Cooperation in Science
and Technology

United States Department of Economic Affairs,
Education and Research (SAFE)
Federal Office for Agriculture (FOAG)

www.fao.org/soils

Functions of soil

1. Food and biomass production

- agricultural, forest, pasture, and wetland

2. Storing/filtration/transformation

- Soil is responsible for the chemical conversions between minerals, organic matter and water.
- Soil diversifies chemical substances during biochemical processes in soil.
- Soil is a natural filtering barrier forming clear underground water sources.

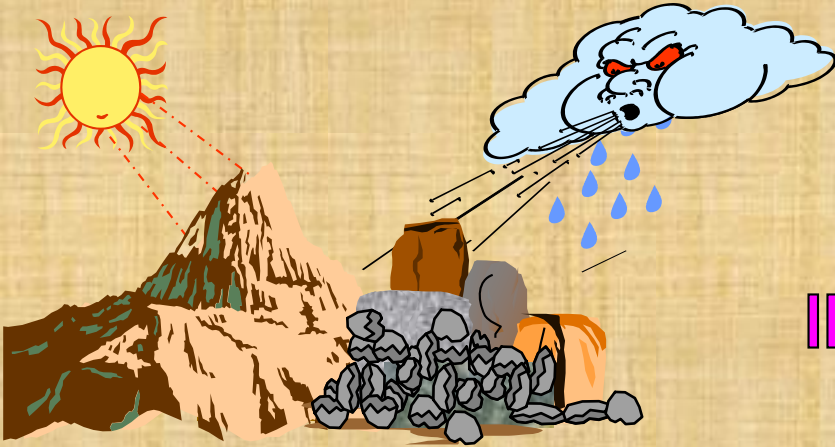
3. Habitate and gene pool:

- Soil hosts very large amount and variety of organisms as a living environment.

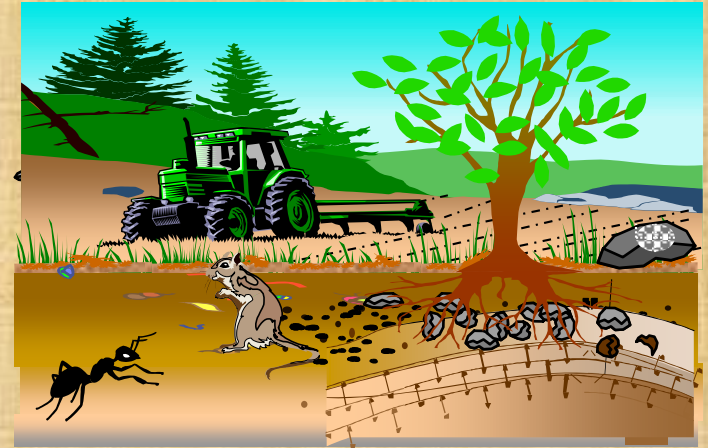
4. Soil as a raw material:

- clay, sand, gravel, mineral and peat,

Soil formation



Physical and chemical weathering



Plants, animals, human



1 cm toprak



200-1000 year
(average 500 yr)



Humus formation

Ideal agricultural soil







Horizons

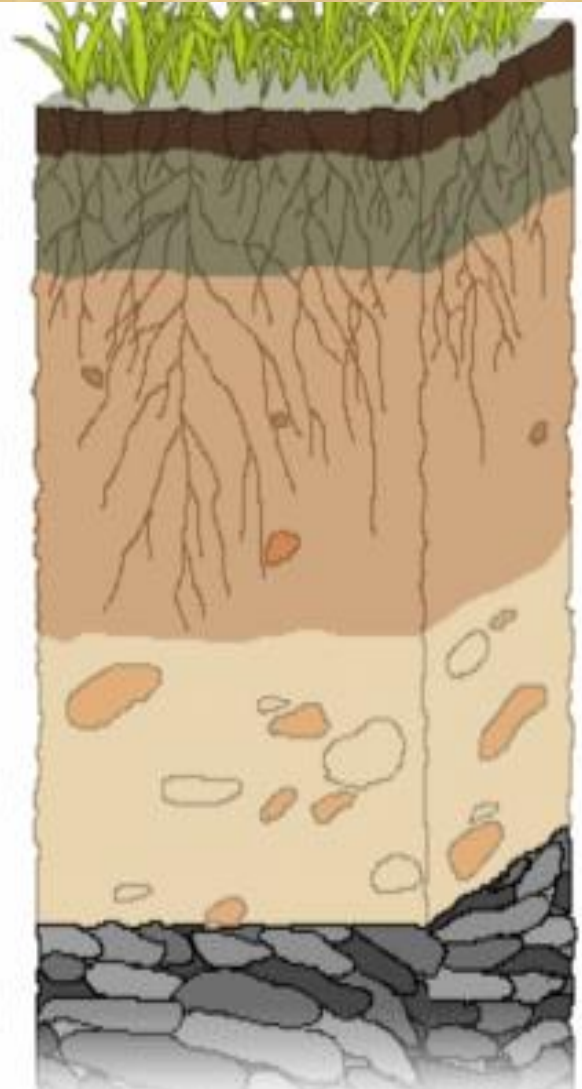
O (Organic)

A (Surface)

B (Subsoil)

C (Substratum)

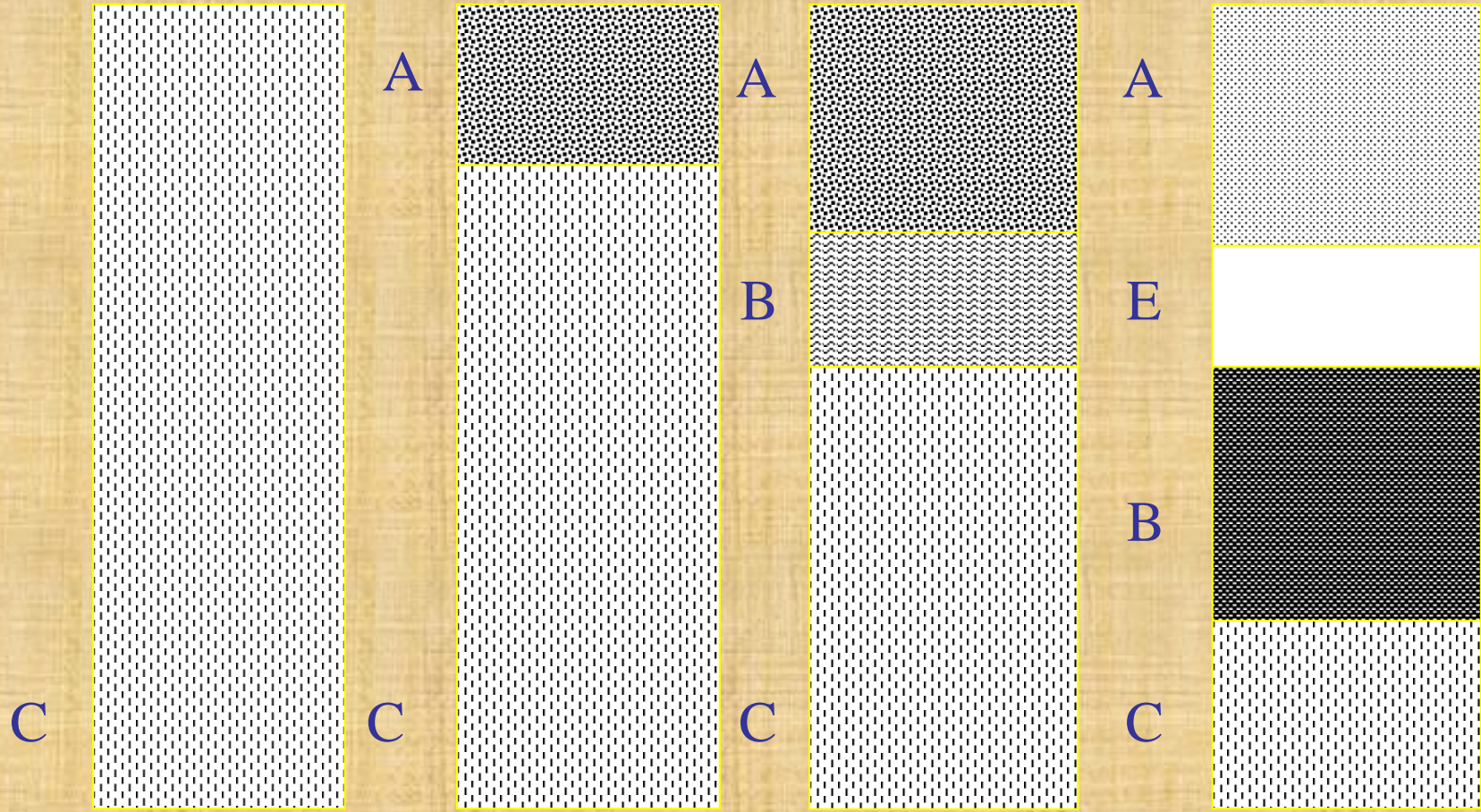
R (Bedrock)



Parent material

Young soil

Mature soil



Soil formation

ELAPSED YEARS : 0 YR

Parent Material existed
in humid and hot region

R

ELAPSED YEARS : 10 YR

Weathering Rocks

WEATHERING describes the means by which soil, rocks and minerals are changed by physical and chemical processes into other soil components.

Weathering is an integral part of soil development. Depending on the soil-forming factors in an area, weathering may proceed rapidly over a decade or slowly over millions of years.

C horizon develops on above Regolit.

C

R

ELAPSED YEARS: 100 YR

Vegetation formation,
accumulation of organic matter

Dying plants are accumulated on
Surface and form soil organic matter
and horizon-A

Weathering process continue through
below profile

Parent material (R), changes into
Horizon-C.

A

C

R



ELAPSED YEARS : 1000 YR

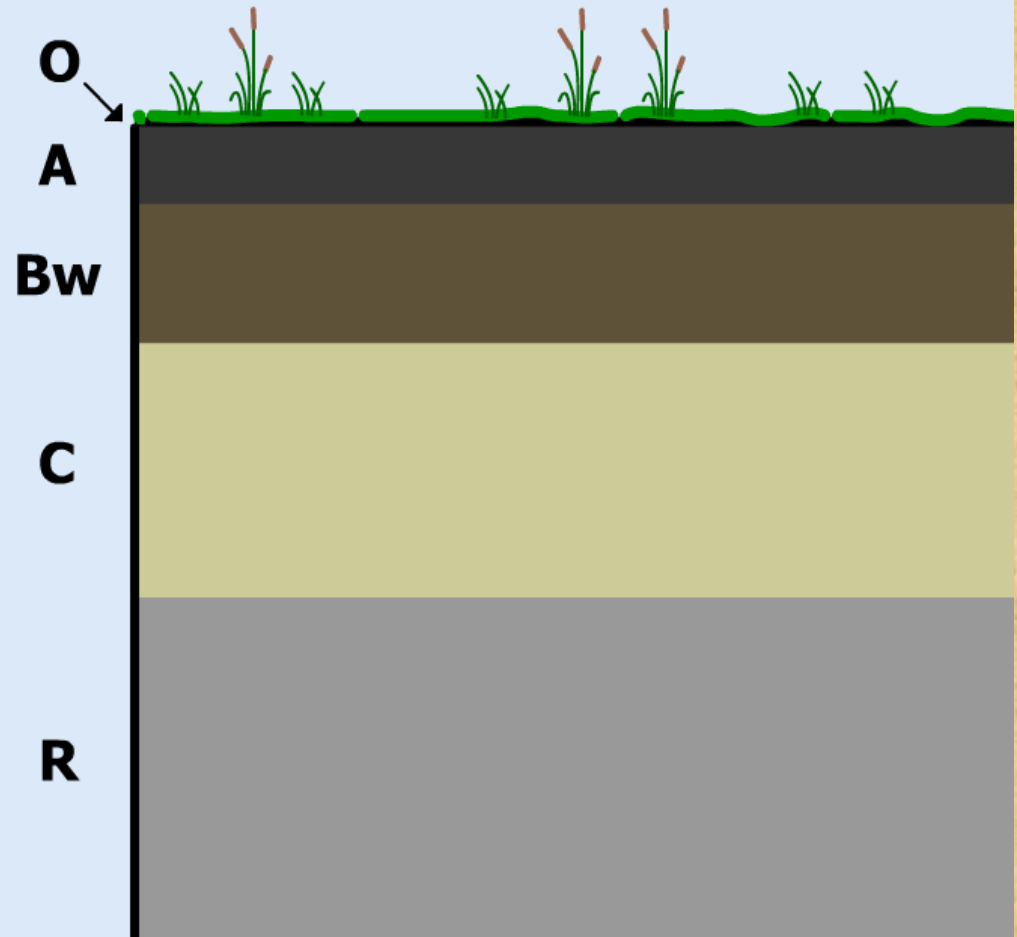
Color and structure
development

Thickening horizon-A becomes
dark color

Horizon-O can form by accumulating
plant debris

Formation of Fe-oxides and clay
minerals on the upper horizon and

Their transport to below layers
causing formation of horizon-Bw



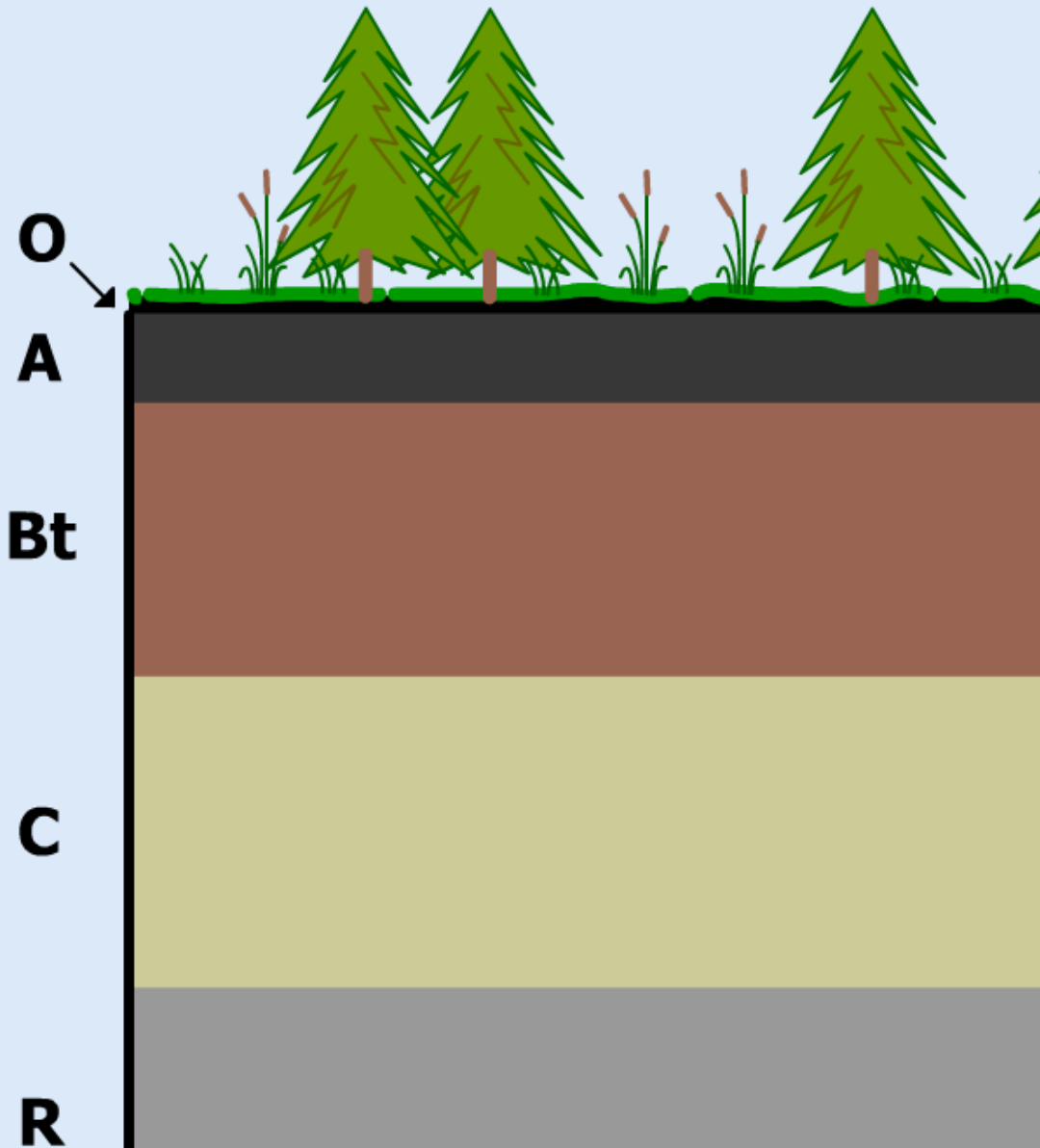
ELAPSED YEARS : 10,000 YR

Increasing clay transport
and accumulation in below layers

Fe-oxides and clays move down and,
Horizon-B becomes more redish
(formation of horizon-Bt

Thickening and darkening horizon-A
by increasing organic matter
accumulation

Weathering process continue through
below layers

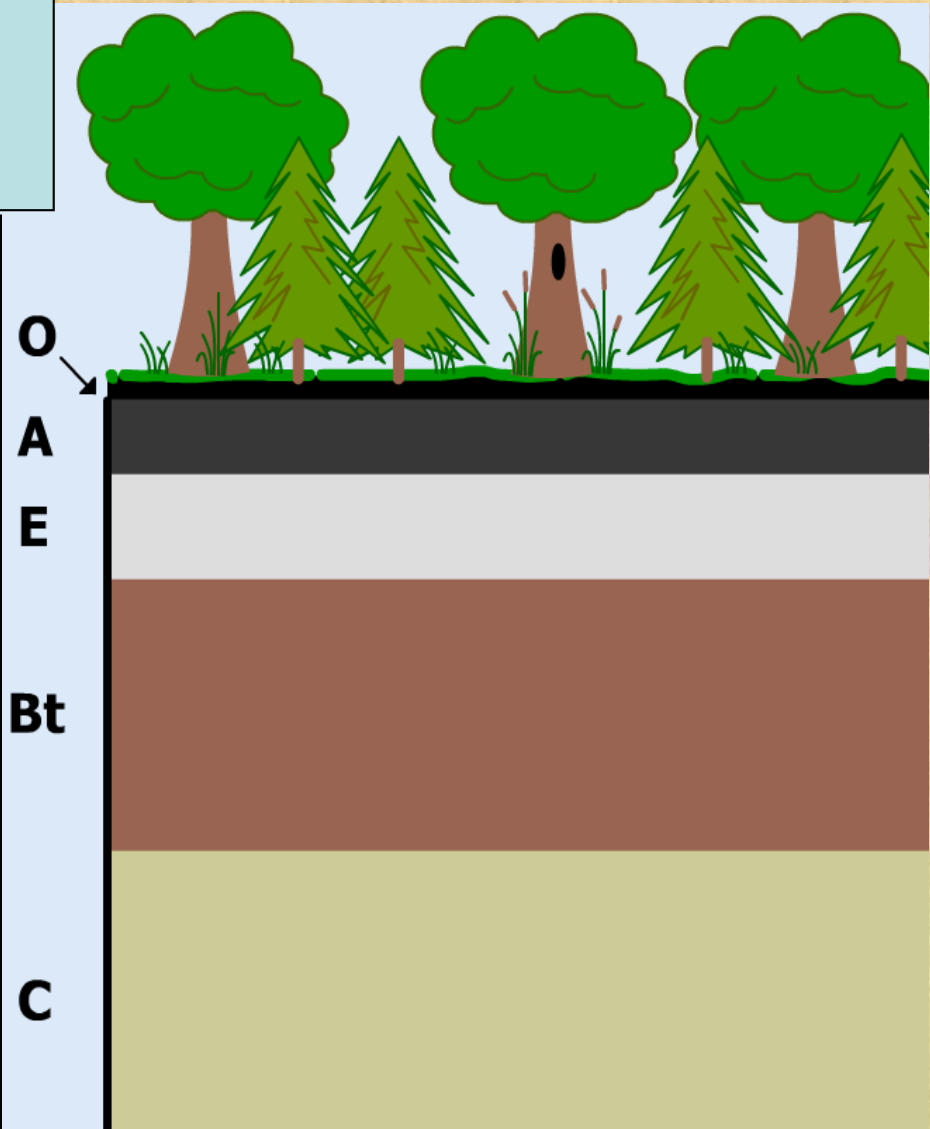


ELAPSED YEARS : 100,000 YR

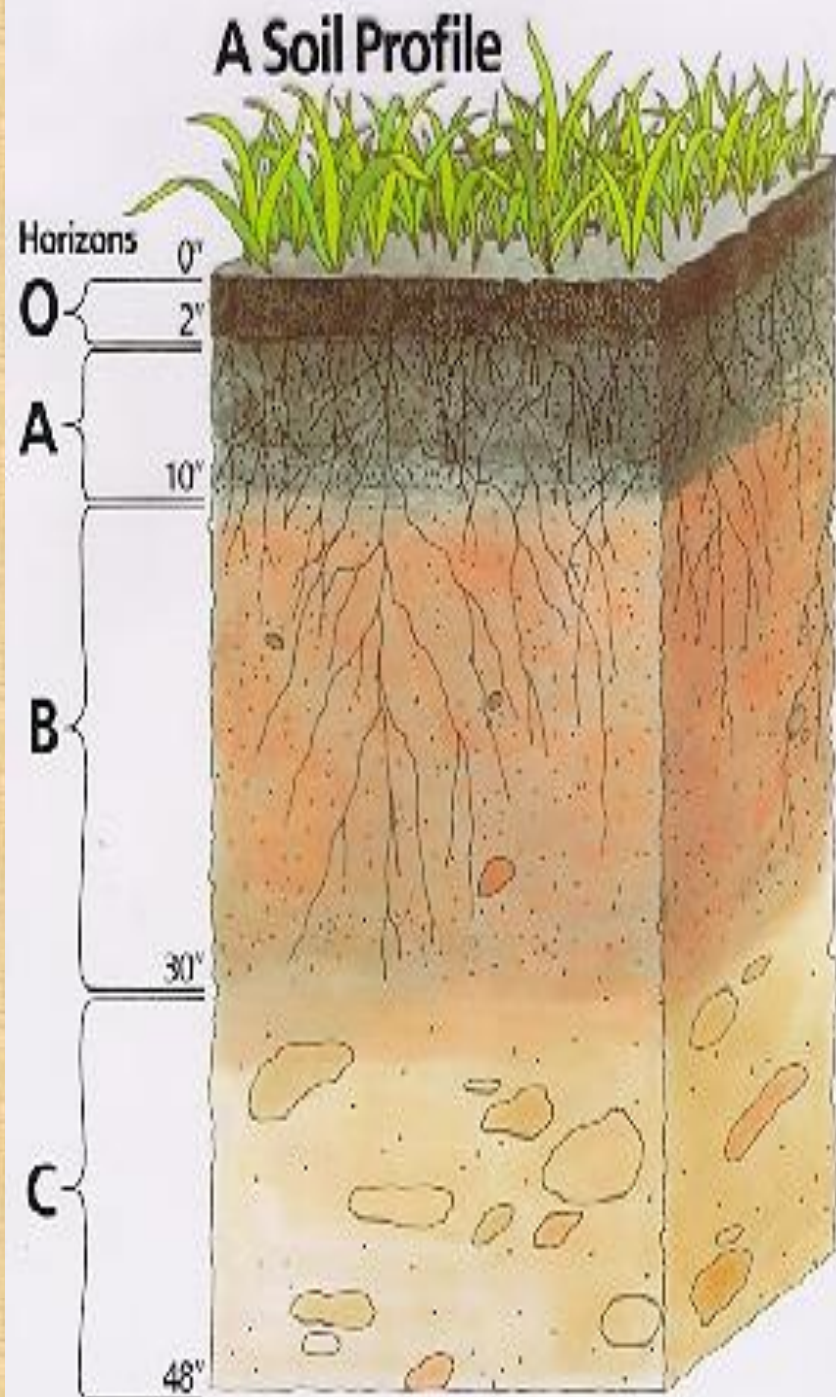
Continous leaching of Fe-oxides and silicate clays in upper layers resulted in the formation of Horizon-E.

Increasing clay transport forms a thick and deep horizon-Bt.

Weathering process continue through below layers



And 100,000
yr later...



Why soil is vital?

1 grams of soil (within 15cm depth)

600.000.000 Bacteria

400.000 Fungi

100.000 Algae

**And many bugs and animals
as well...**

The UNESCO Courier / june1997



Soil forming factors...

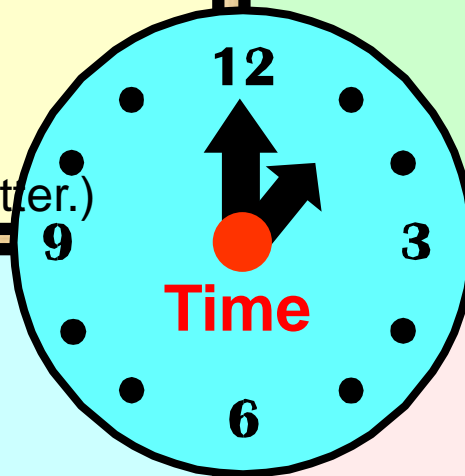


Parent material

(Rocks, minerals, organic matter.)

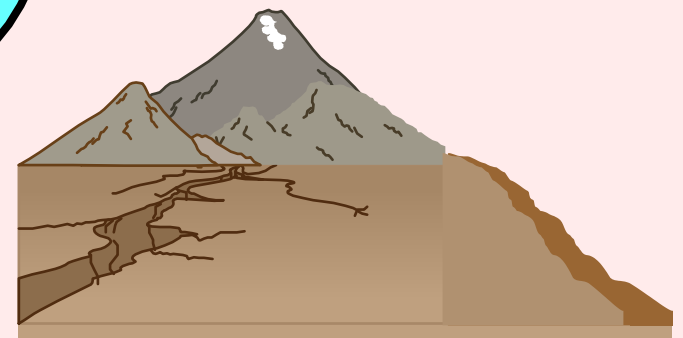


Climate (Rainfall, temperature and wind)



Organizms

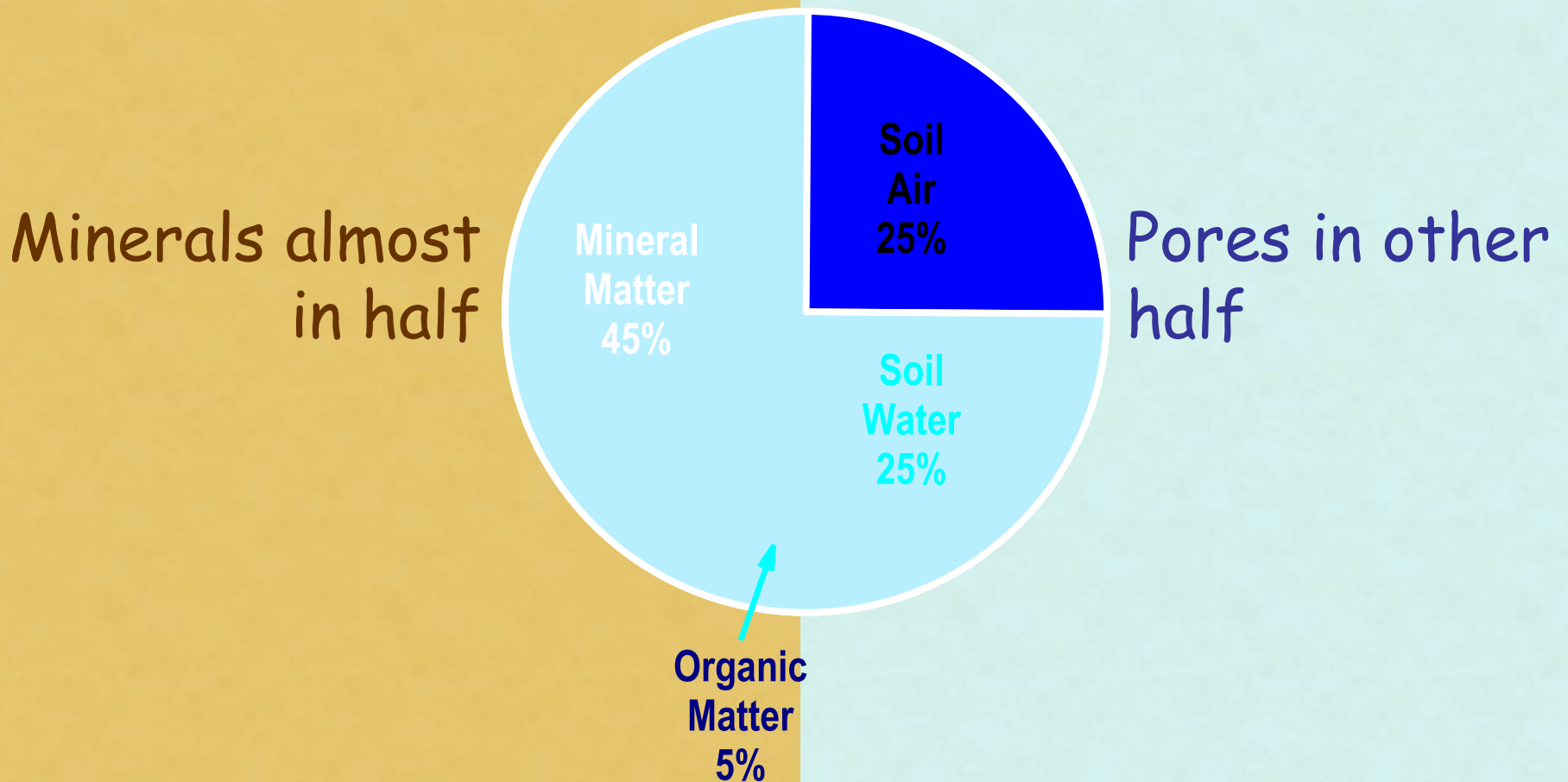
(Plants and animals)



Topography

(Altitude, slope, vector)

SOIL COMPONENTS (4 divisions)



PHYSICAL CHARACTERISTICS OF SOIL

- Color
- Texture
- Structure
- Pores
- Soil water
- Soil air

<https://www.qld.gov.au/environment/land/soil/soil-properties/colour/>

Why different color?



Why different pattern (structure)





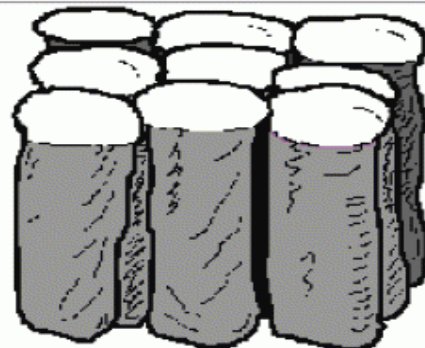
Granular: Resembles cookie crumbs and is usually less than 0.5 cm in diameter. Commonly found in surface horizons where roots have been growing.



Blocky: Irregular blocks that are usually 1.5 - 5.0 cm in diameter.



Prismatic: Vertical columns of soil that might be a number of cm long. Usually found in lower horizons.



Columnar: Vertical columns of soil that have a salt "cap" at the top. Found in soils of arid climates.



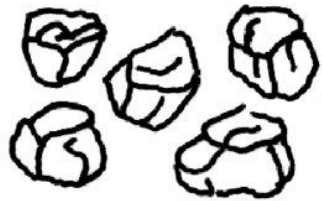
Platy: Thin, flat plates of soil that lie horizontally. Usually found in compacted soil.



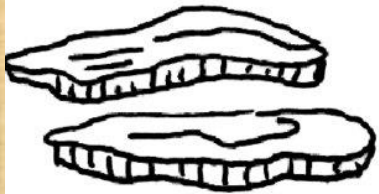
Single Grained: Soil is broken into individual particles that do not stick together. Always accompanies a loose consistence. Commonly found in sandy soils.



Prismatic



Structureless:
massive



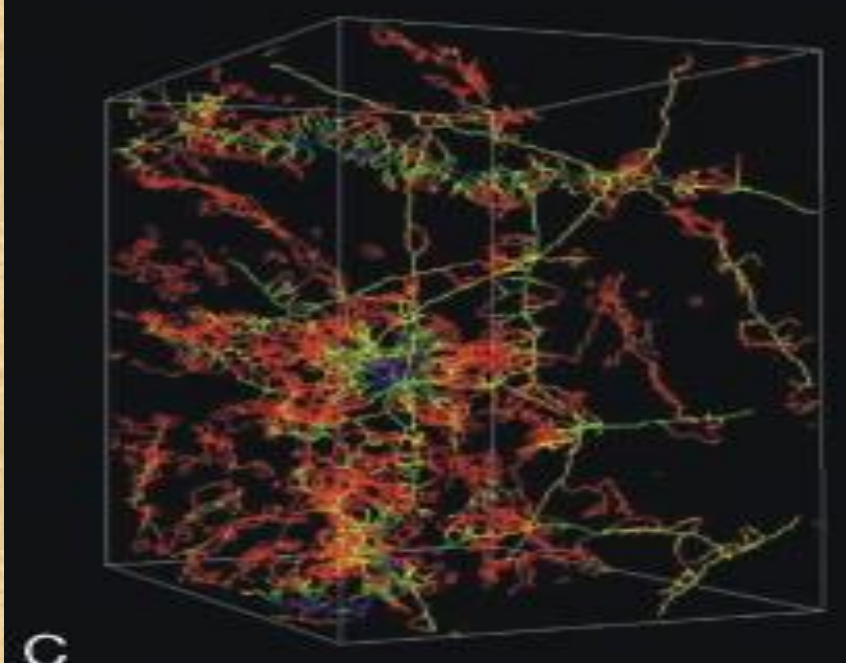
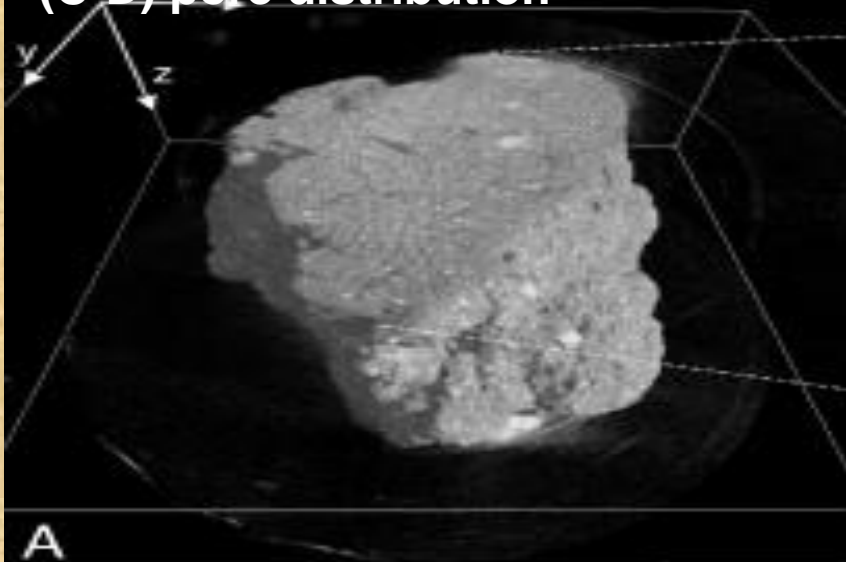
Structureless:
single grain



Space in soil (porosity)

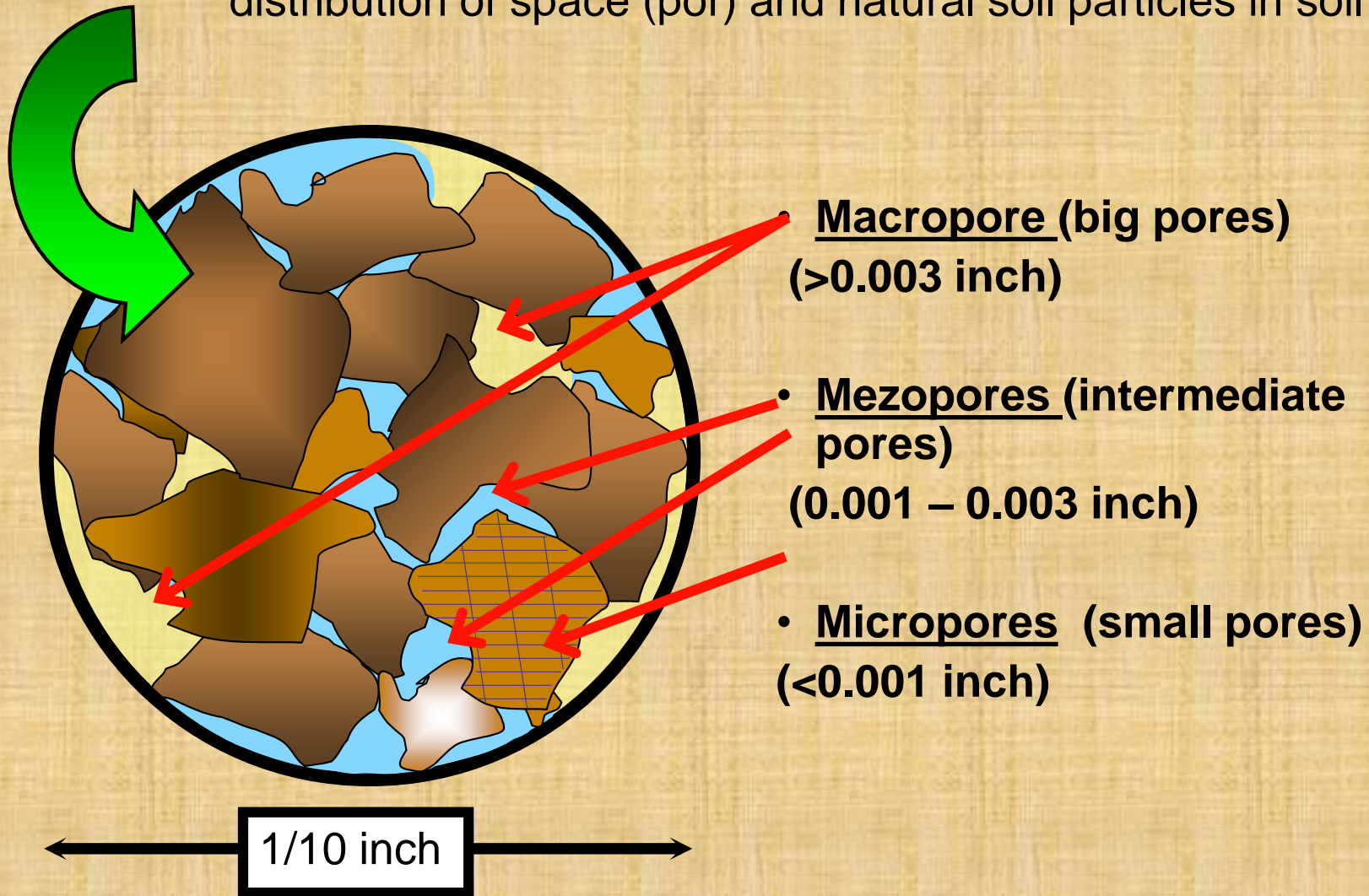


(A) A soil particle with 5mm dimension in 3D, (B) cross section view, (C-D) pore distribution

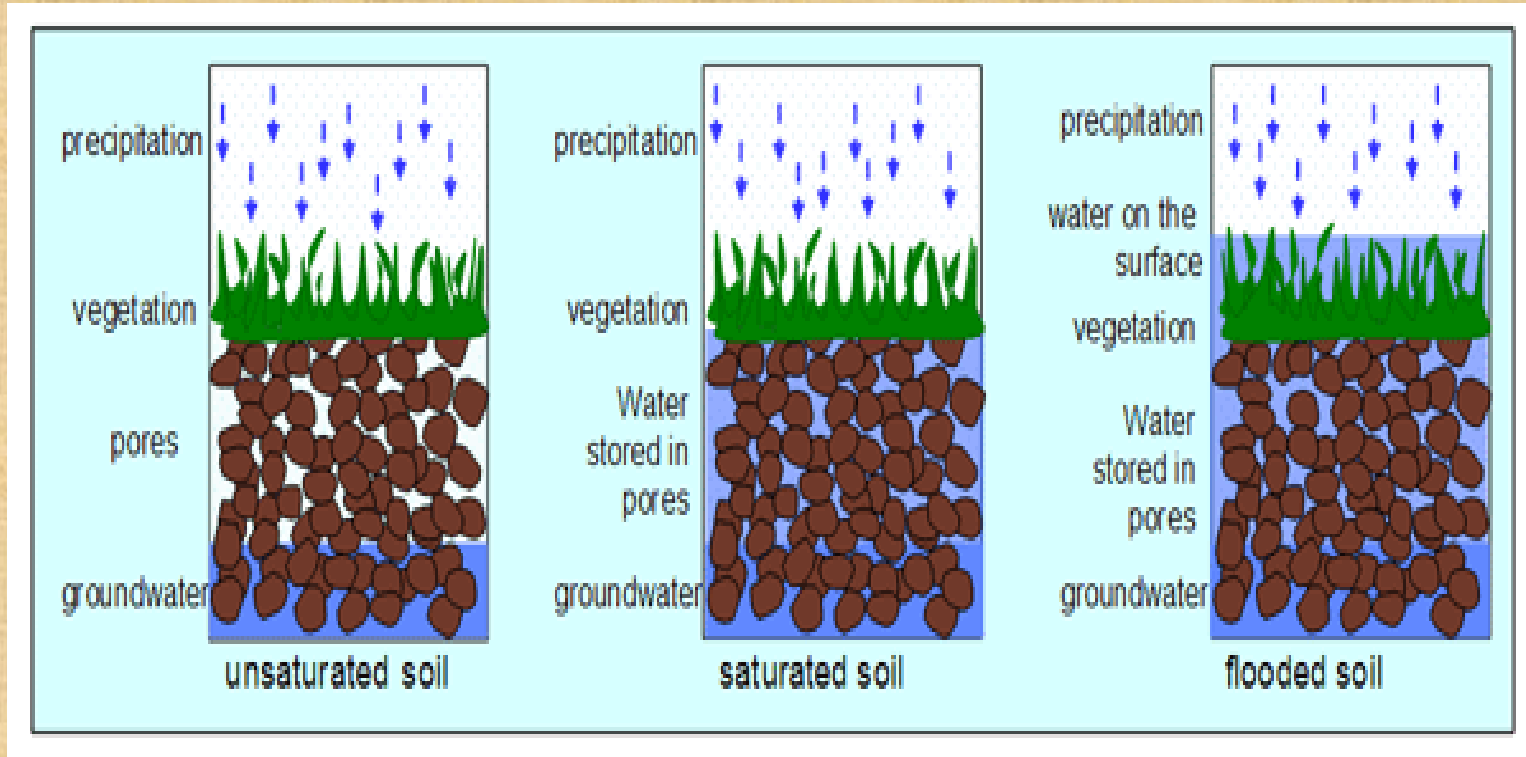


Soil porosity

distribution of space (por) and natural soil particles in soil



Changing water in soil



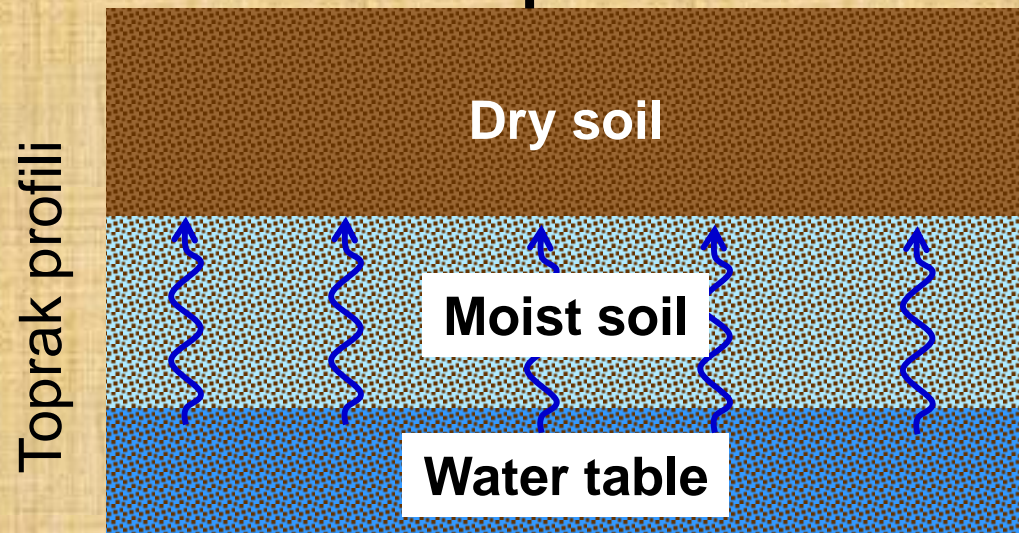
Water in soil

Two main forces moving water in soil

gravity

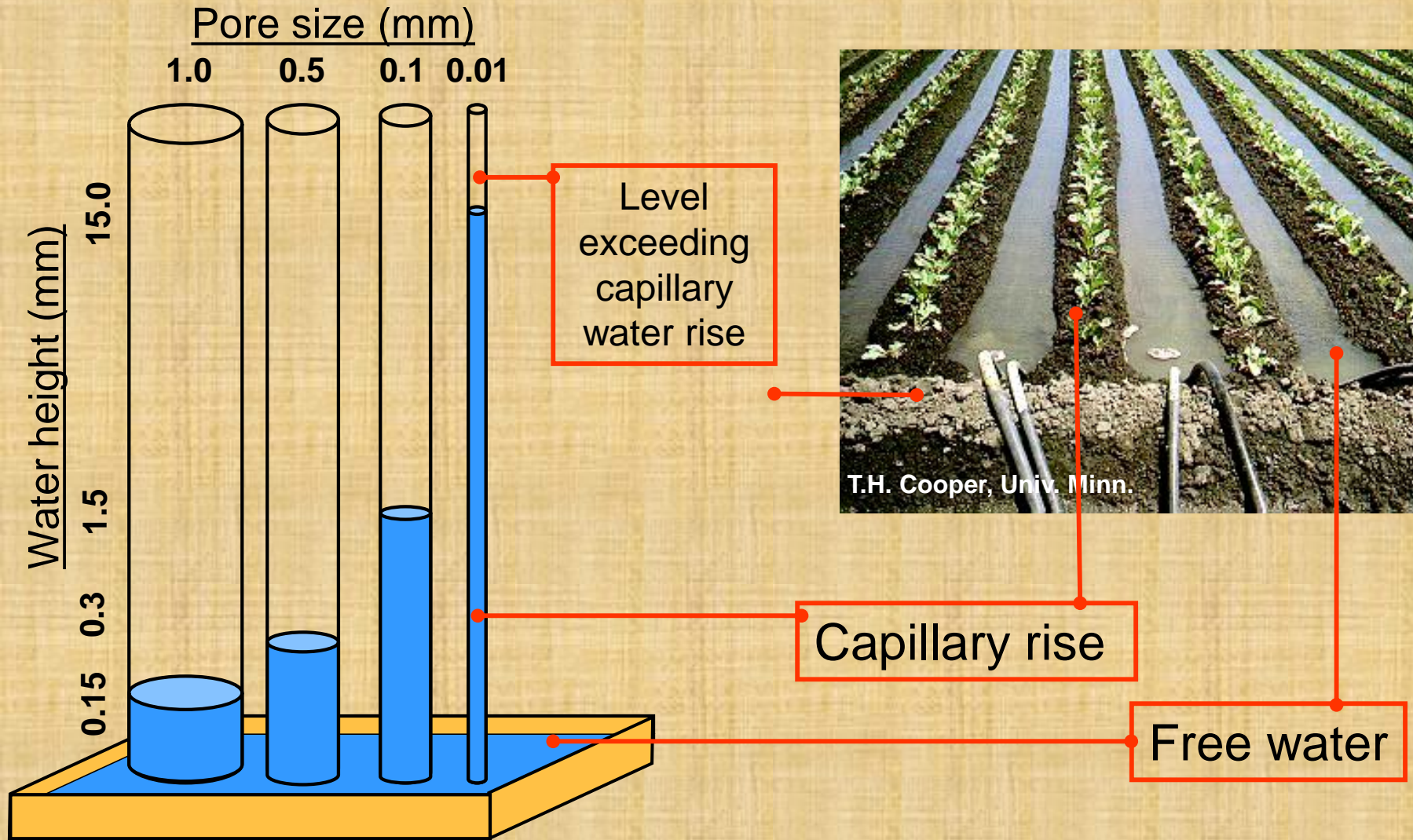


Matric potential



In conditions with less water in the micropores, physical attraction force between soil-water surface is driven by slow water movement called "capillarity"

Capillary water movement in soil

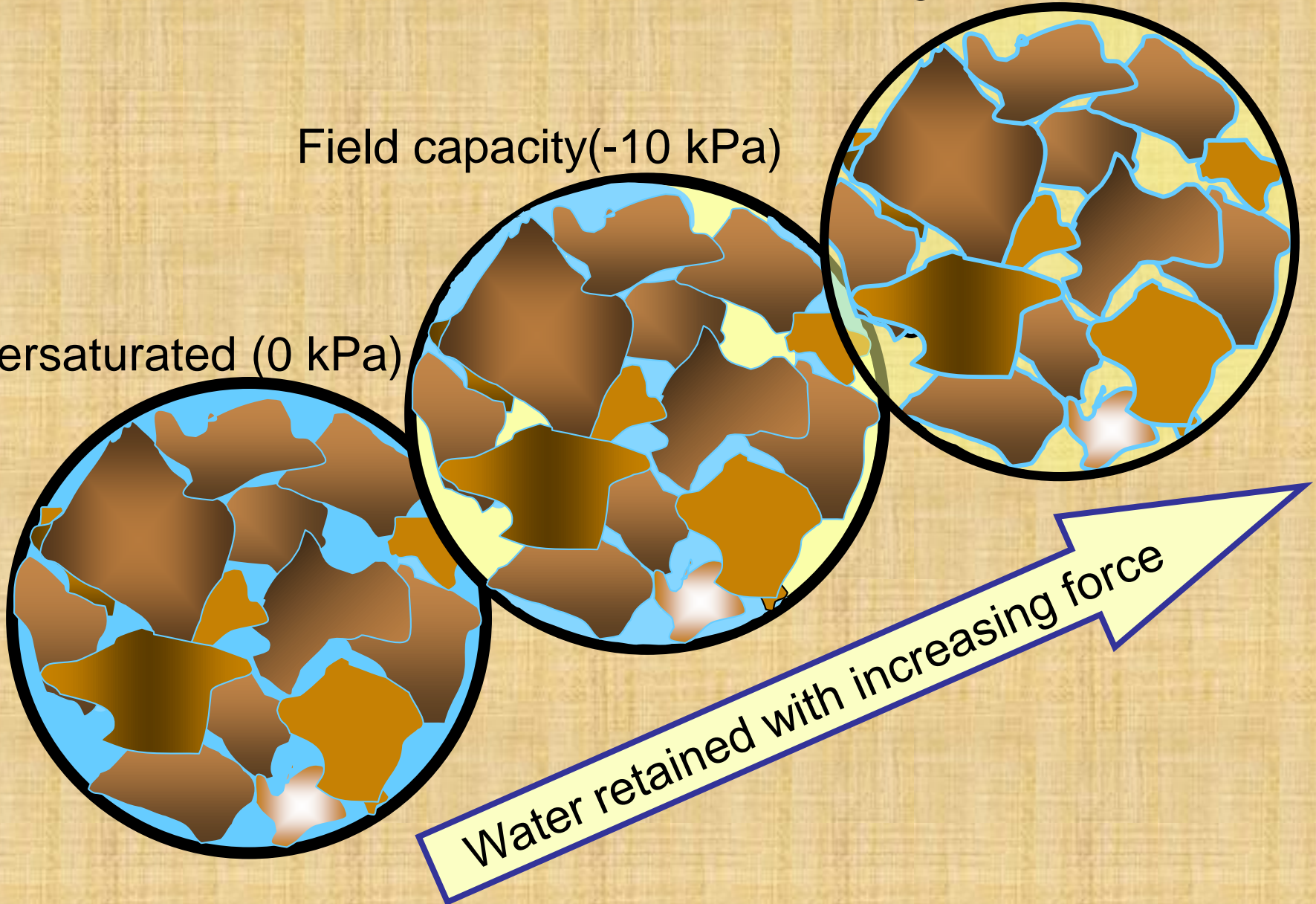


Water retention in soil

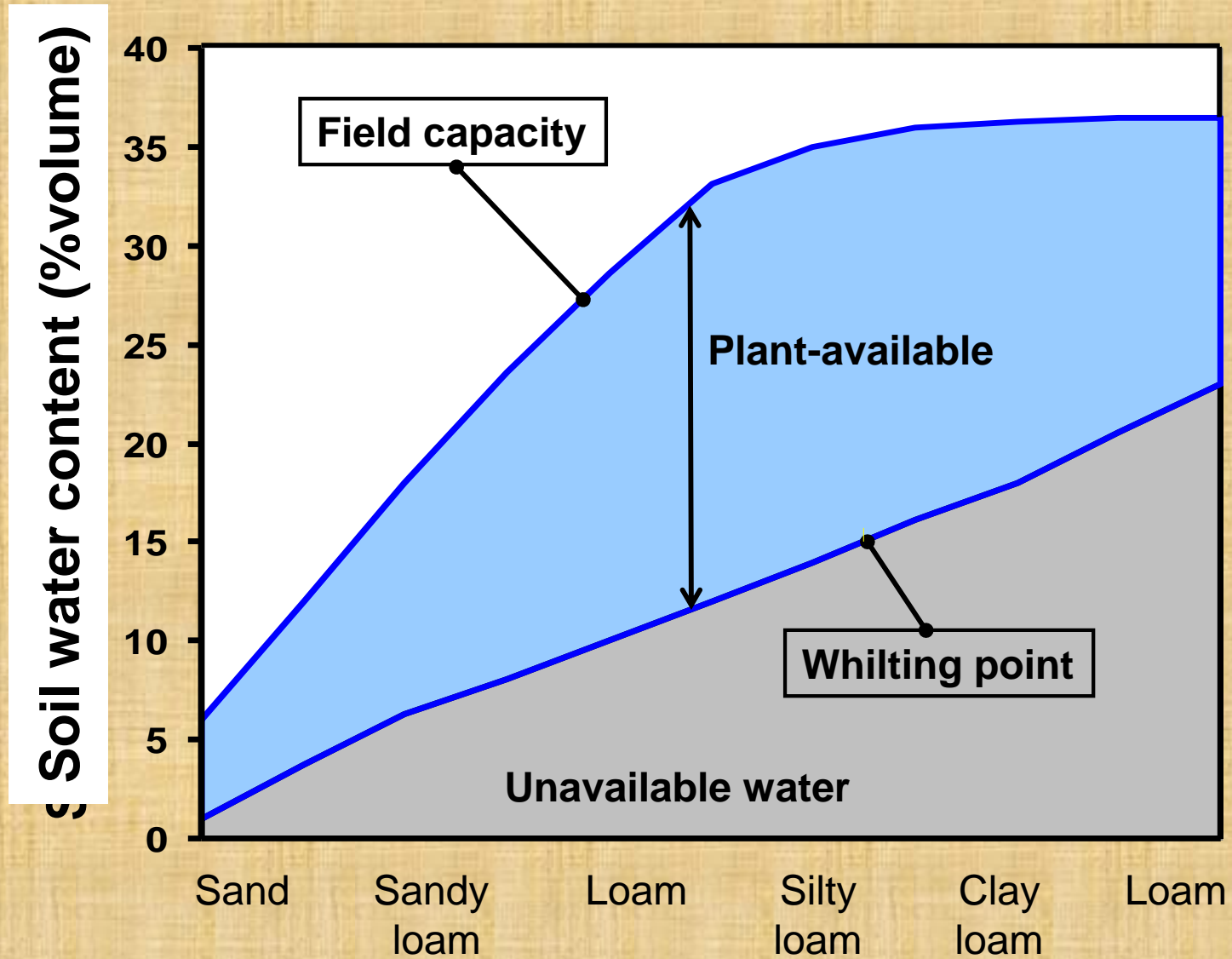
Wilting point (-1500 kPa)

Field capacity (-10 kPa)

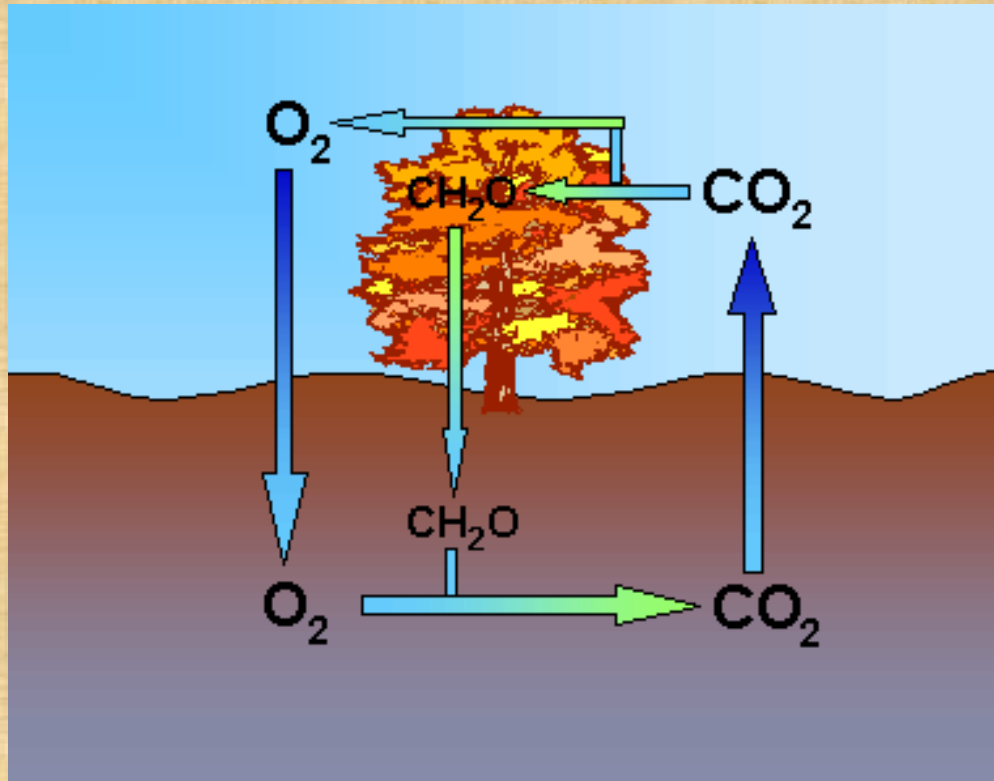
oversaturated (0 kPa)



Plant available water



Air cycle in soil



Soil texture

- Soil mineral portion consists of 'clay', 'sand' and 'silt'



- A ratio between these particles (in %) determines soil texture
 - Coarse texture (more sand-less clay)
 - Heavy texture (more clay –less sand)
 - Loamy(same amount of sand-clay-silt)

Chemical characteristics of soil

- Soil pH
- Mineral nutrients in soil
- Colloids (clay and organic matter)
- CEC (cation exchange capacity)
- Electrical conductivity
- Soil buffering capacity

Soil pH

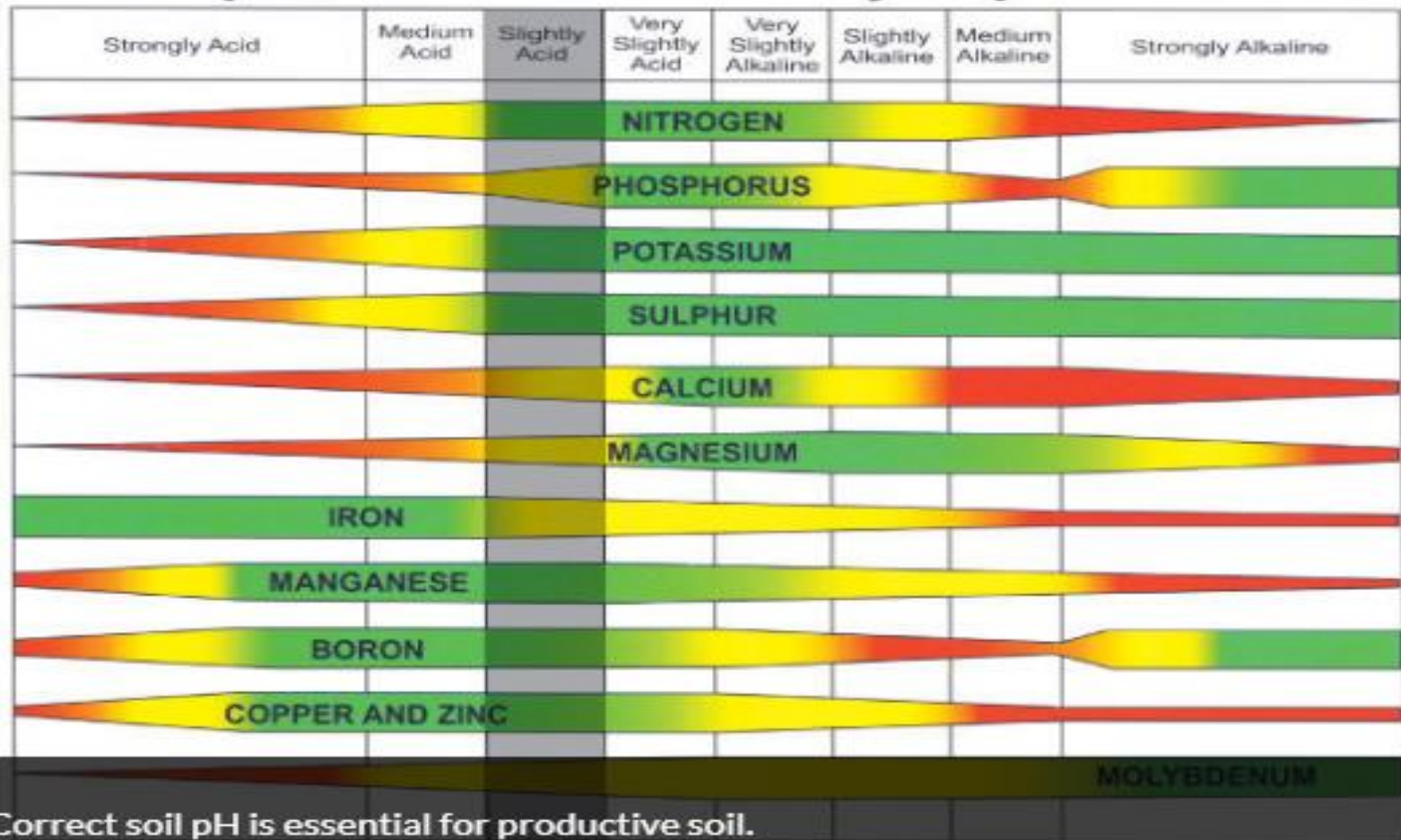
- Soil pH or soil reaction is an indication of the acidity or alkalinity of soil and is measured in pH units.
- Soil pH is defined as the negative logarithm of the hydrogen ion concentration.
- The pH scale goes from 0 to 14 with pH 7 as the neutral point.
- As the amount of hydrogen ions in the soil increases the soil pH decreases thus becoming more acidic.
- From pH 7 to 0 the soil is increasingly more acidic and from pH 7 to 14 the soil is increasingly more alkaline or basic.

What can be the potential reasons for the situation below?



Why soil pH matters??

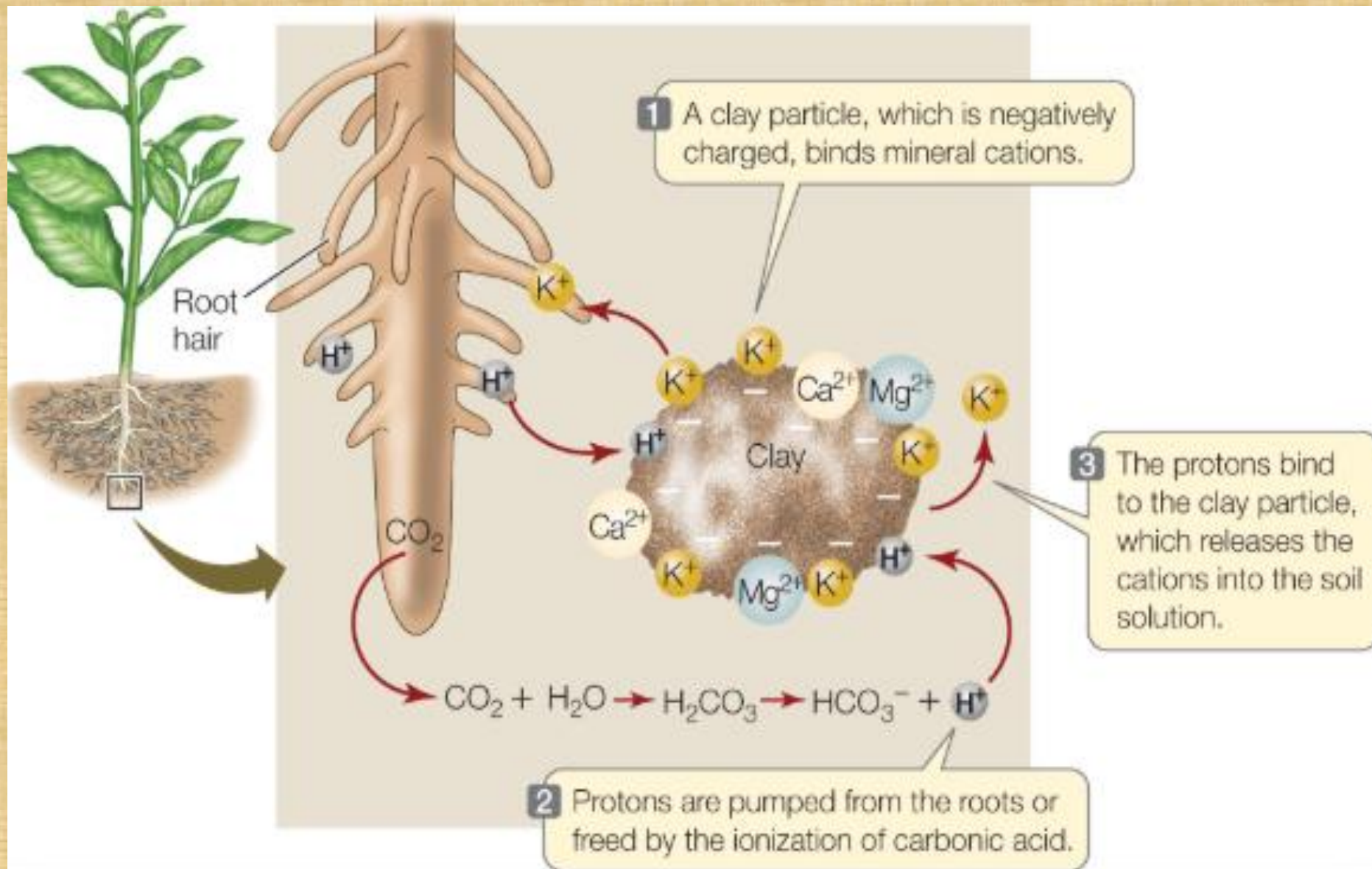
How soil pH affects availability of plant nutrients



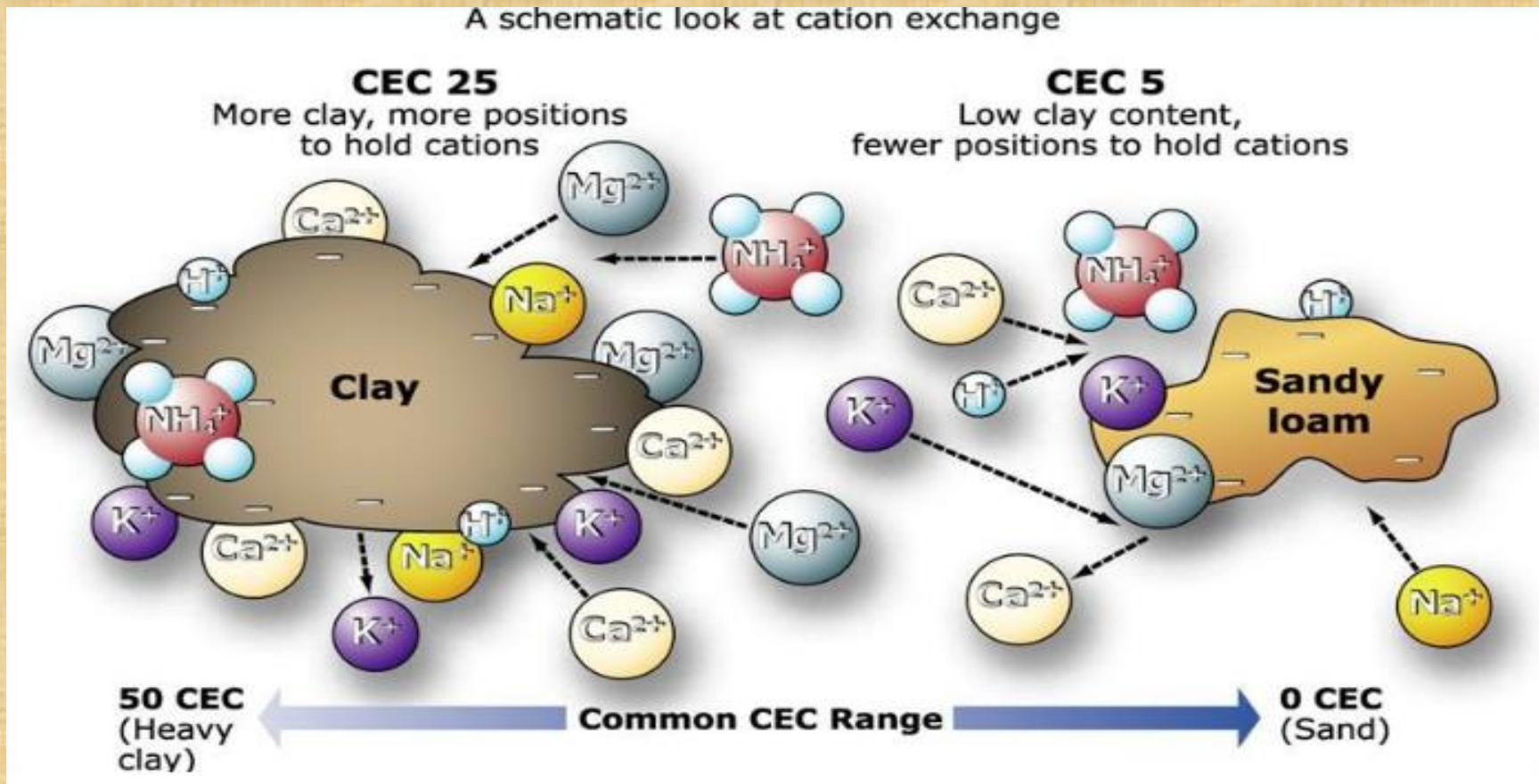
Correct soil pH is essential for productive soil.

4.0 4.5 5.0 5.5 6.0 6.5 7.0 7.5 8.0 8.5 9.0 9.5 10.0

Ion Exchange in Soil



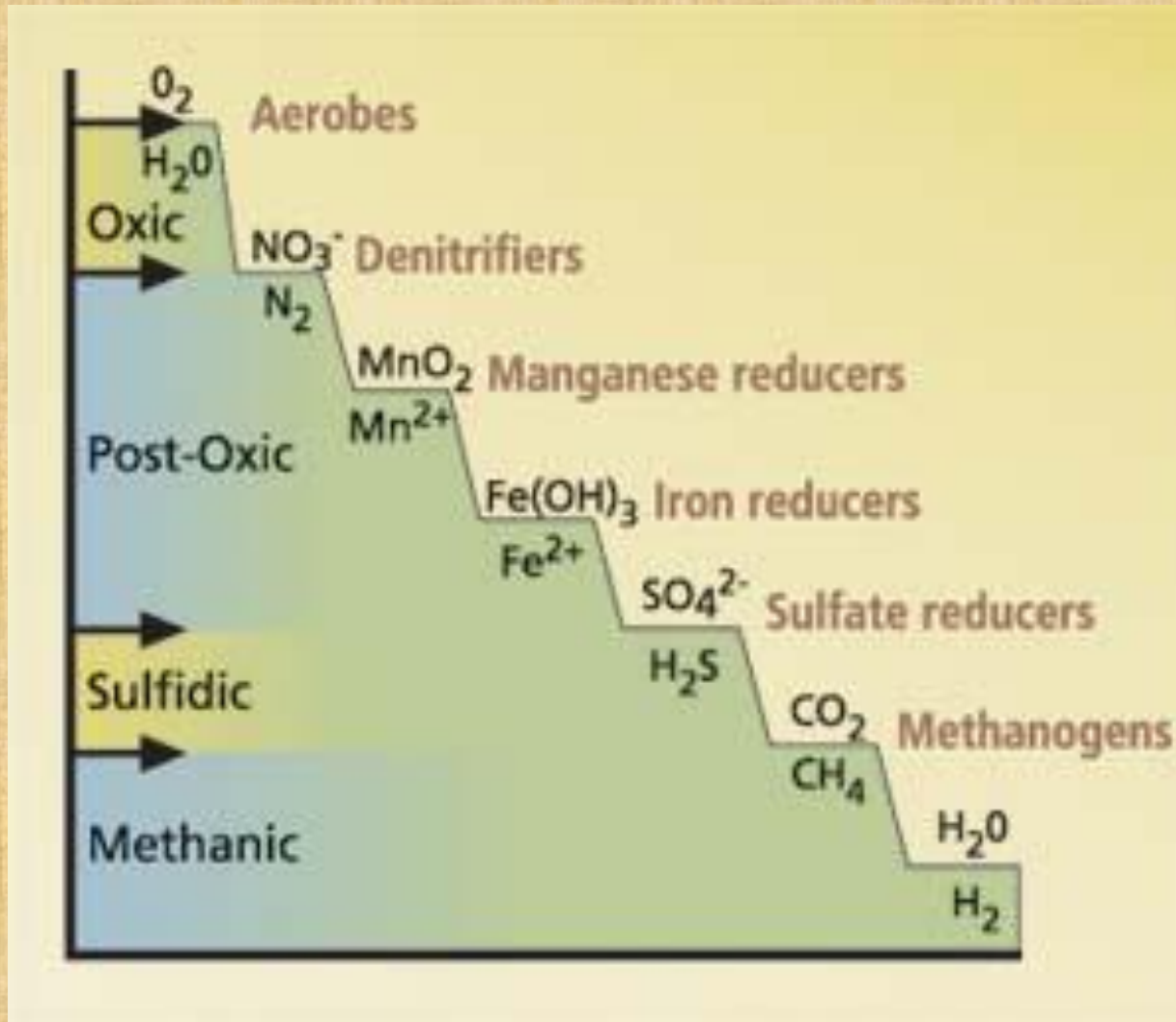
Cation Exchange Capacity (CEC)



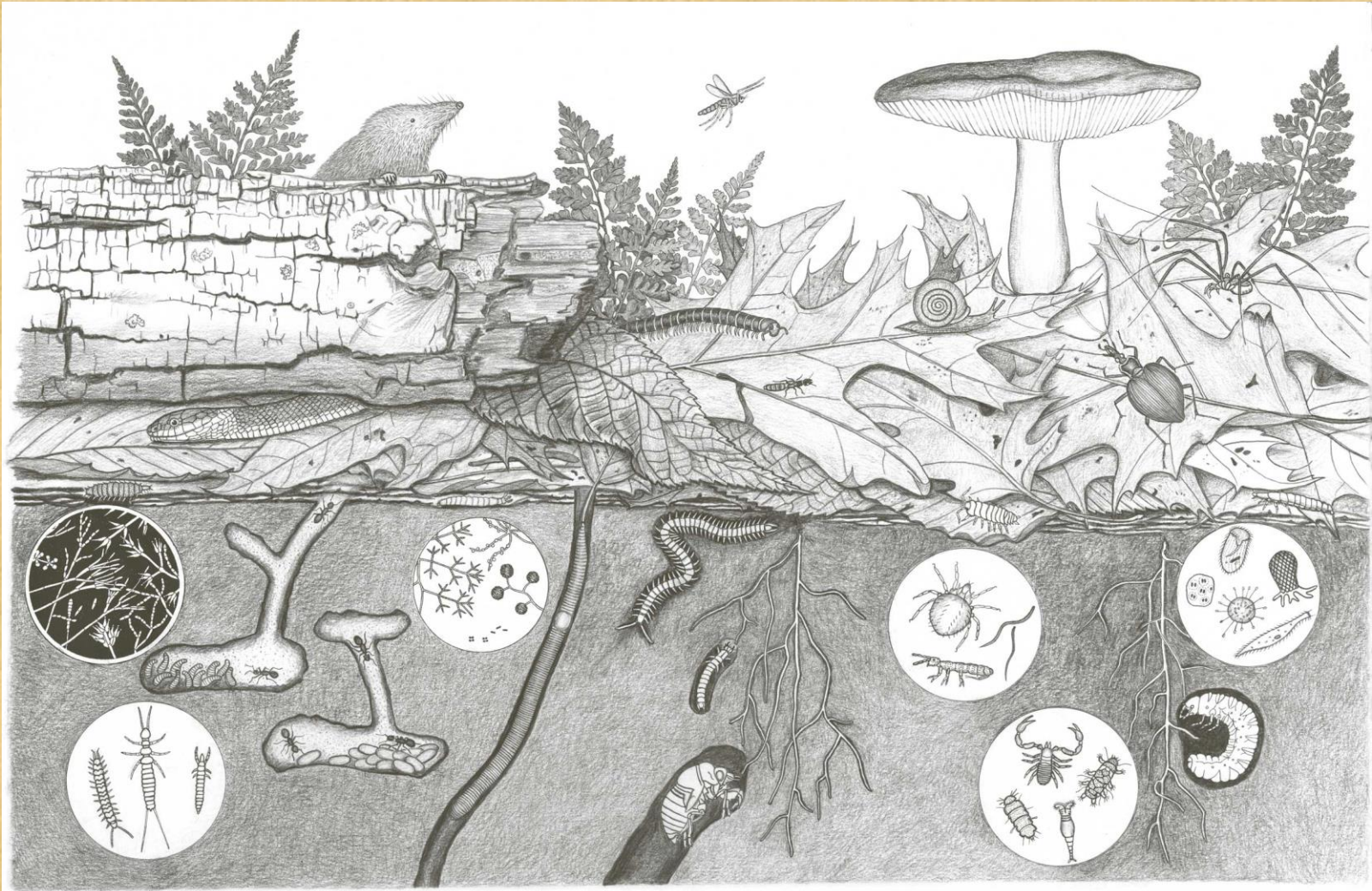
Not only nutrients but also pollutants can be sorbed (adsorption+adsorption) by clays

CEC is important in the sense of soil pollution

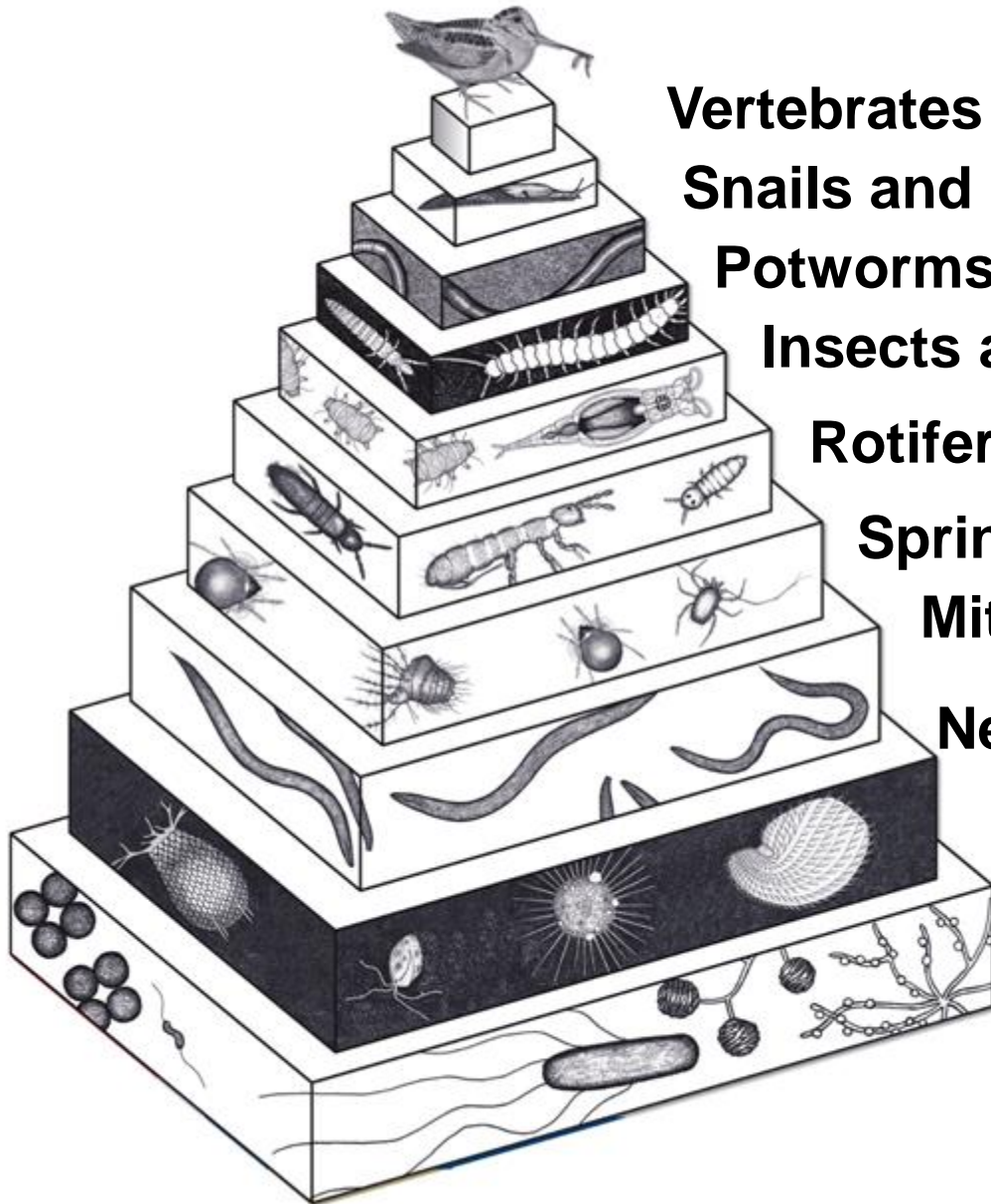
Soil redox potential (oxidation and reduction processes)



Biological characteristics of soil



Biology Pyramid in Soil



Vertebrates (1)

Snails and Slugs (100)

Potworms and Earthworms (3,000)

Insects and Spiders (5,000)

Rotifers (10,000)

Springtails (50,000)

Mites (100,000)

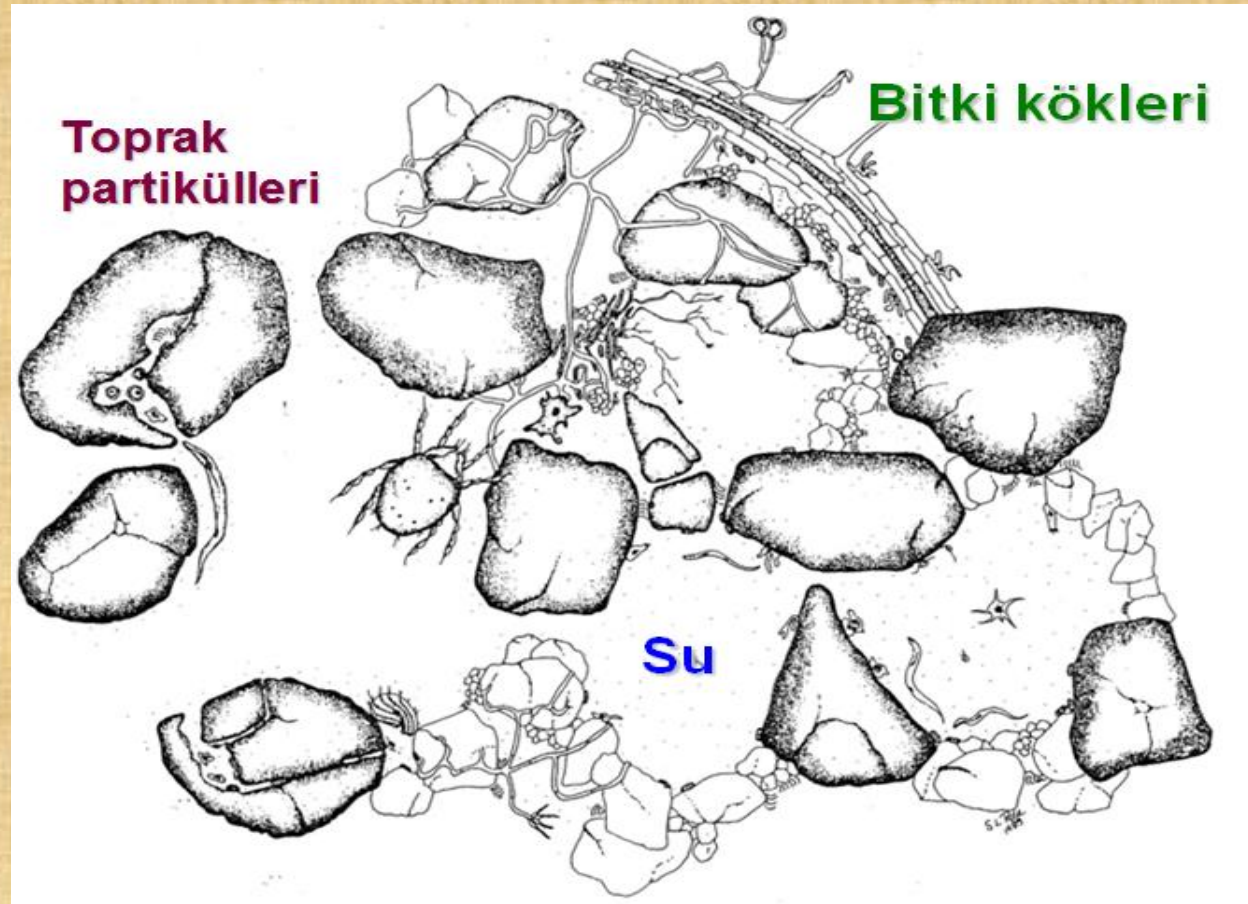
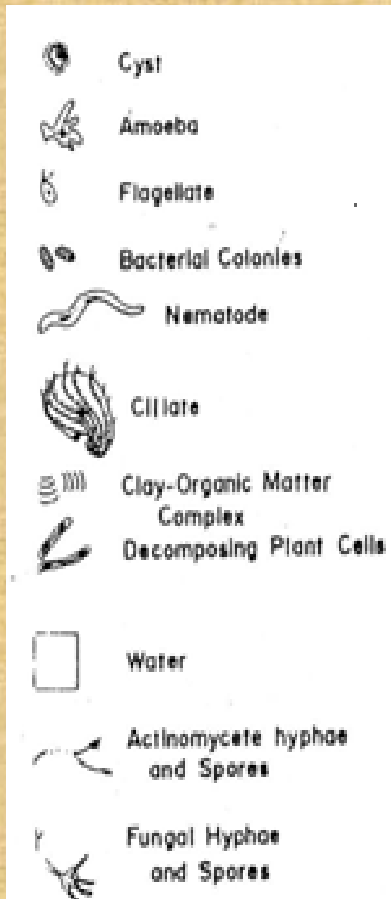
Nematodes (5,000,000)

Protozoa (10,000,000,000)

Fungi (100,000,000,000)

Bacteria (10,000,000,000,000)

Soil as a habitate



1cm² of soil microzone indicating relationships between living organisms and their access to life resources

Classification of Microorganisms (according to their size)

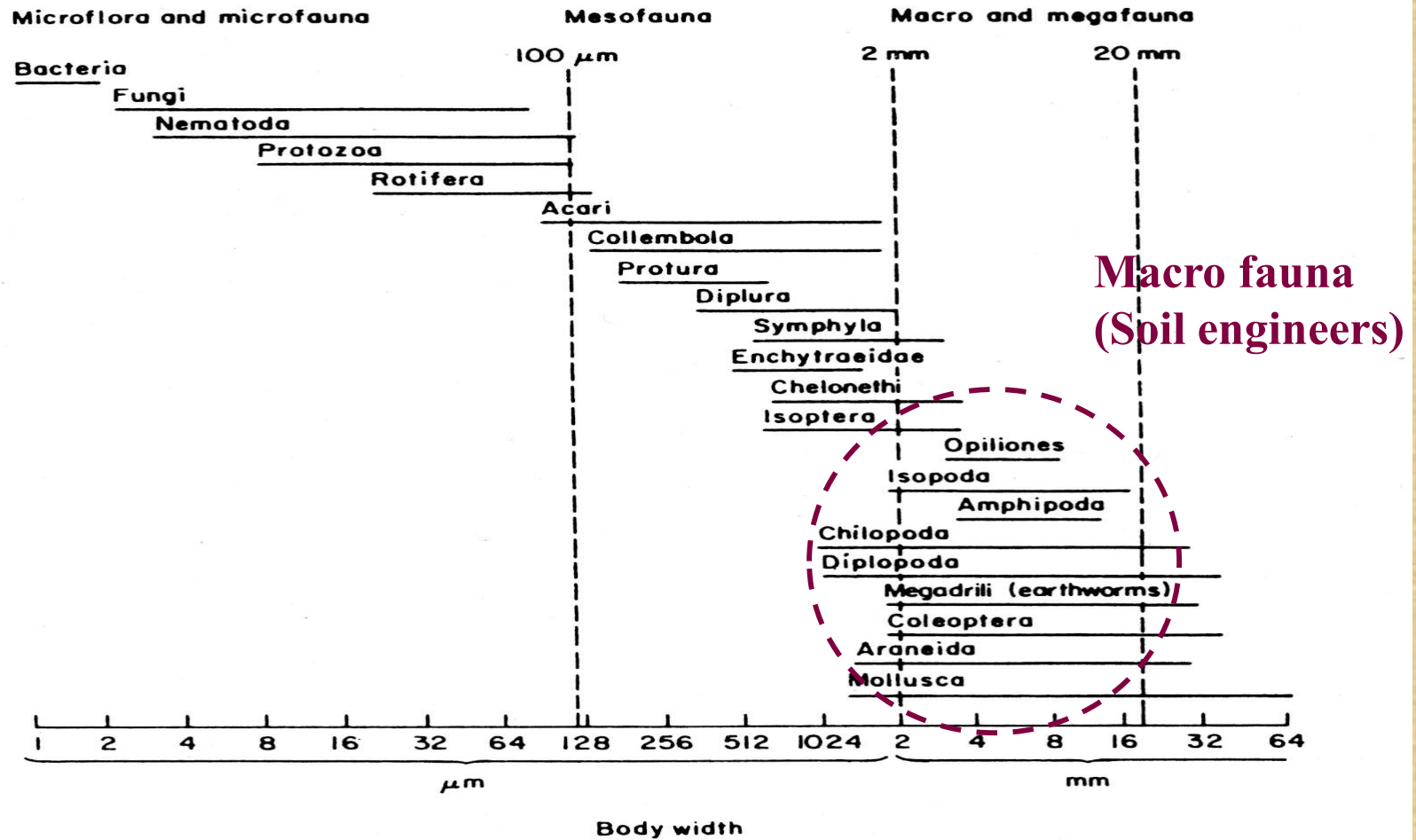


FIGURE 4.3 Size classification of organisms in decomposer food webs by body width (Swift *et al.*, 1979).

Soil Macrofauna

Termit



Örümceğimsiler
(*Pseudoscorpion*)



Solucan



Çiyan (*Centipede*)



Salyangoz (*Snail*)



Macrofauna is important because they're responsible for organic matter decomposition, predation and bioturbation (mixing of mineral soil by living microorganisms)

Mesofauna:
predators
pathogenes

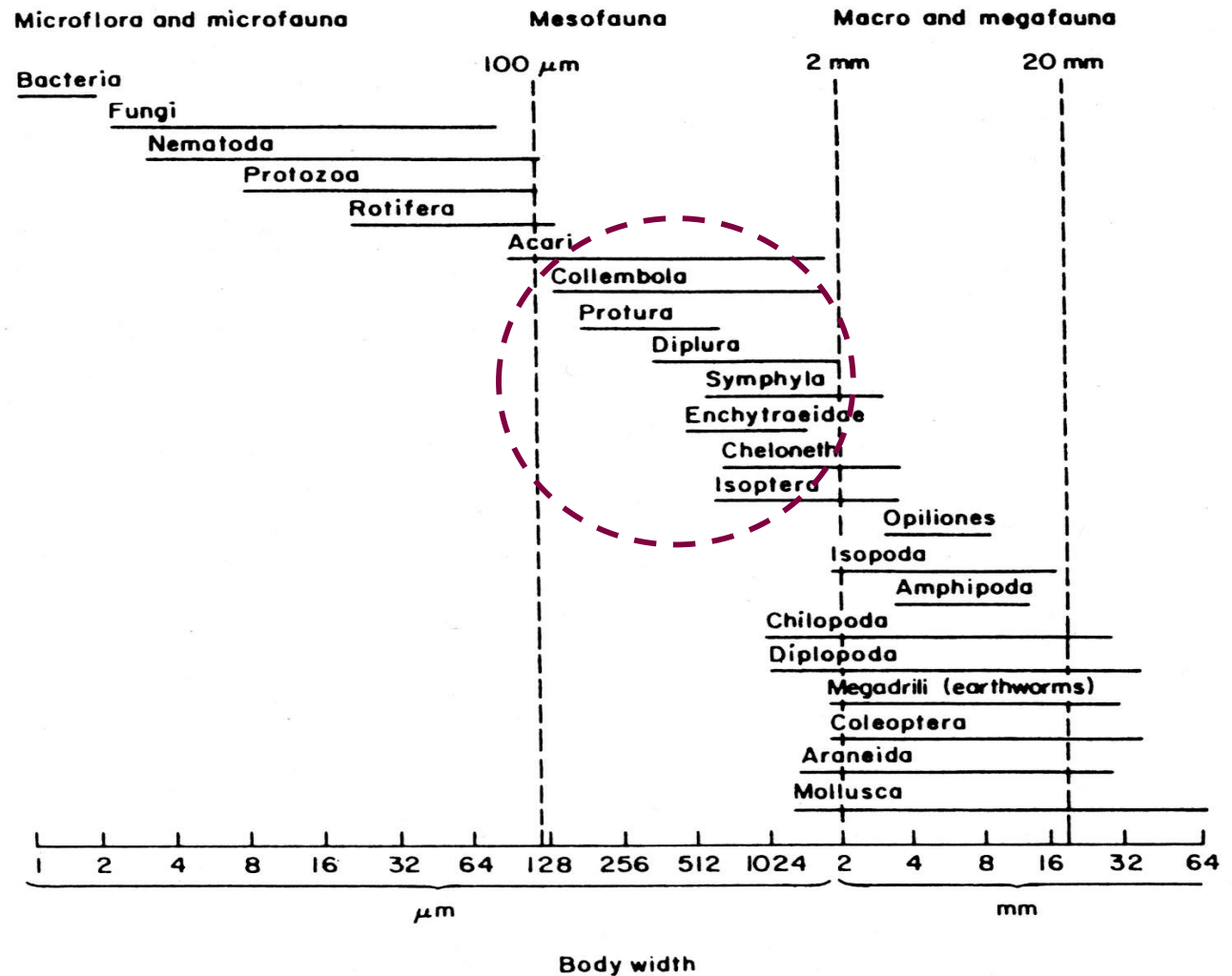
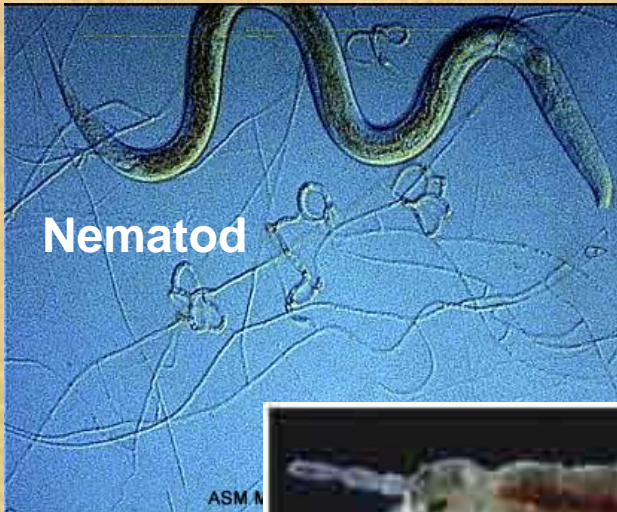
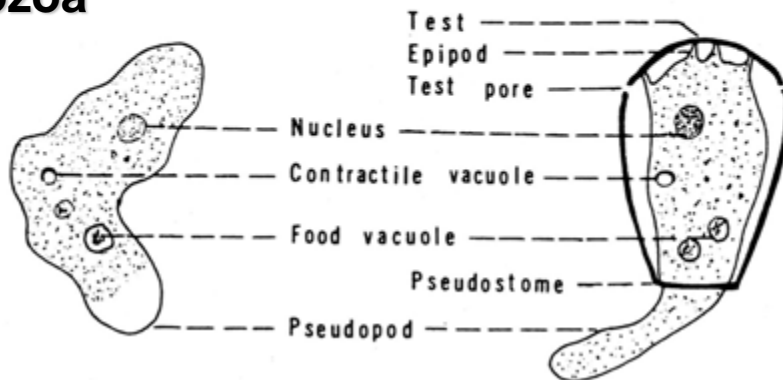
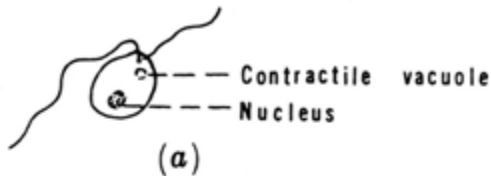


FIGURE 4.3 Size classification of organisms in decomposer food webs by body width (Swift *et al.*, 1979).

Soil Mesofauna



Protozoa



Mesofauna is important because responsible for organic matter decomposition, predation and controlling pathogenes in soils

**Microorganisms
(responsible for
many nutrient
processes in the soil)**

C, N, S, P

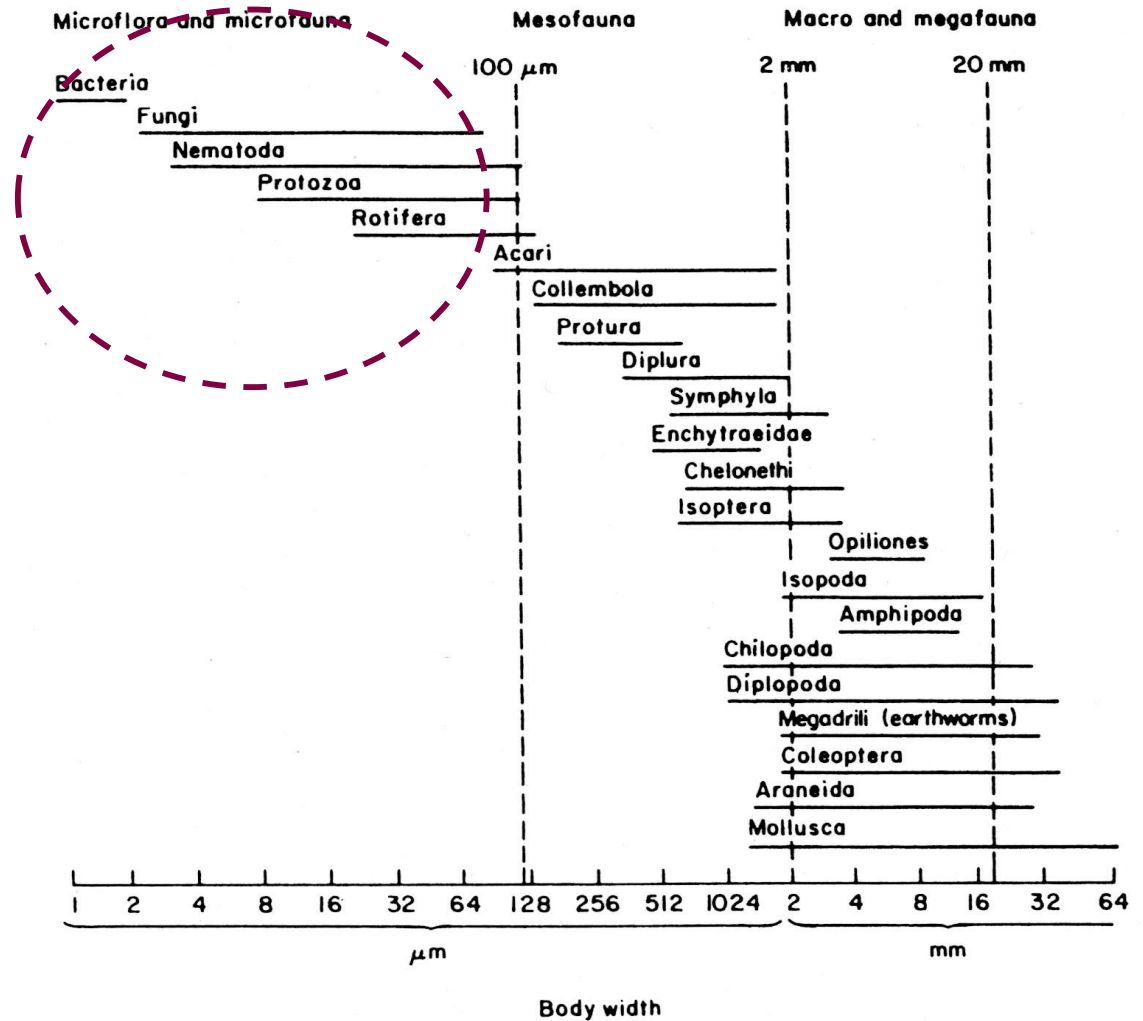
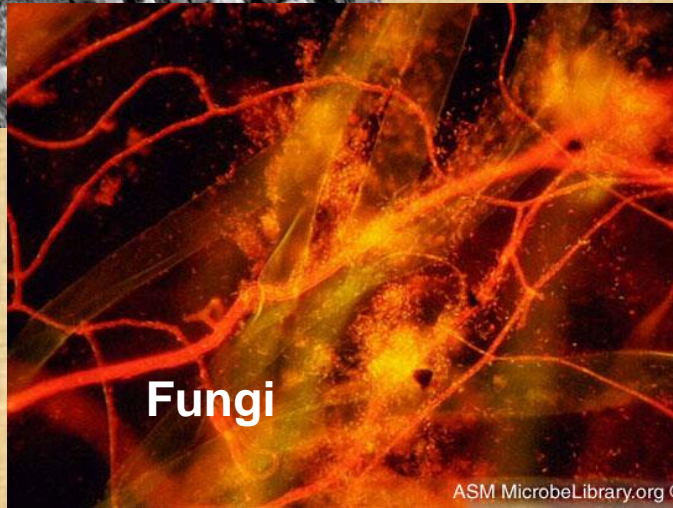
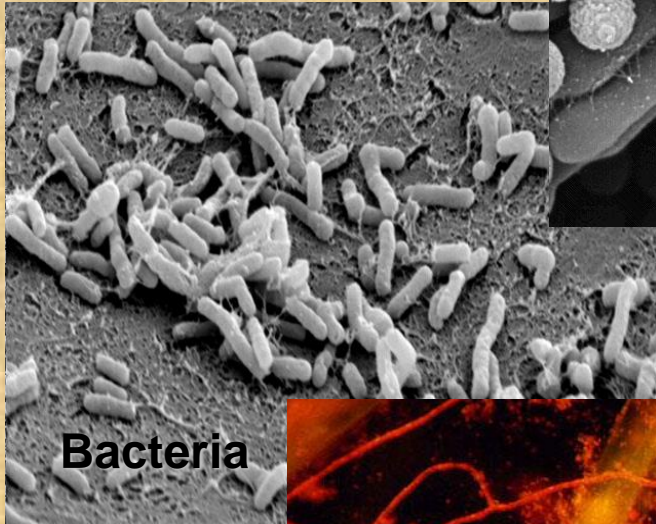


FIGURE 4.3 Size classification of organisms in decomposer food webs by body width (Swift *et al.*, 1979).

Soil microfauna



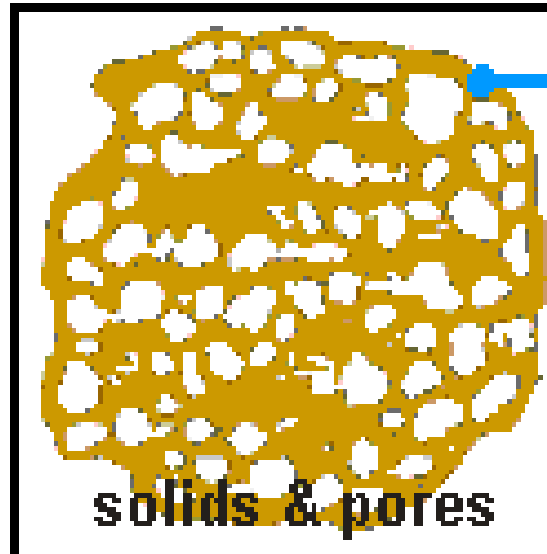
Microfauna is the most abundant in the soil and mainly responsible for the decay of organic matter, nutrient transformations and cycles, “carbon sequestration” and disease suppression as well . Therefore it is regarded as living part of soil organic matter.

SOIL AGGREGATION

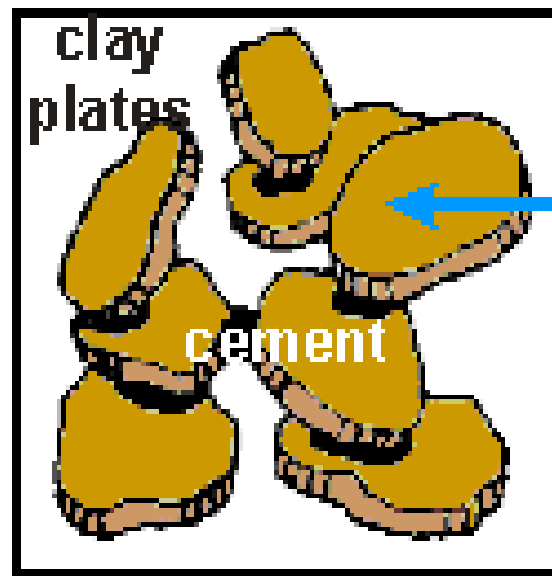
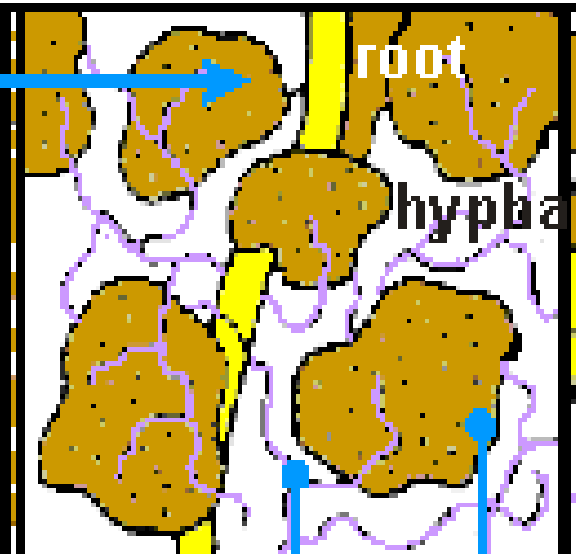
Magnification shows the ever finer structure of soil. Five steps of 10x

(after Tisdall & Oades, 1982)

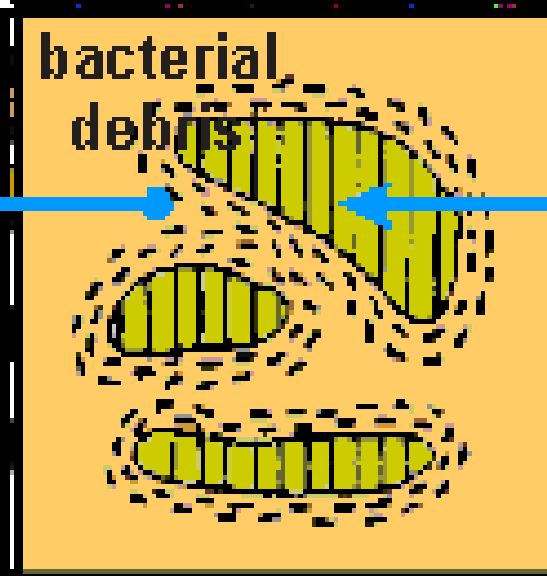
2mm crumb



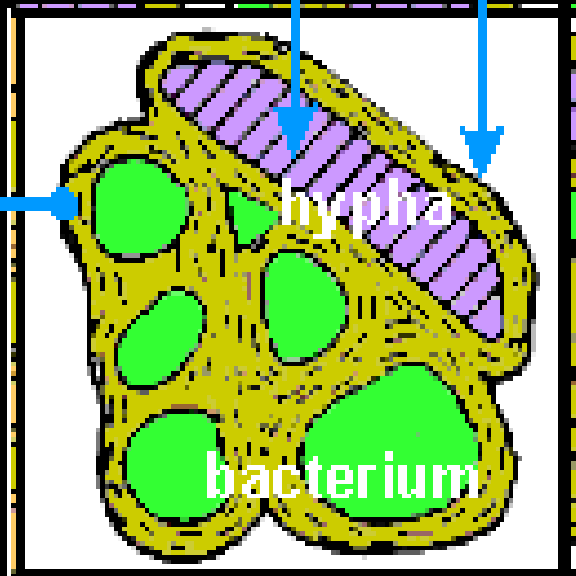
0.2mm particle



200nm clay particle



2000nm clay packet



0.02mm aggregate

Microbial fungi decomposition (nutrient cycle, humus formation, carbon sequestration etc)



FUNGI
1,000x

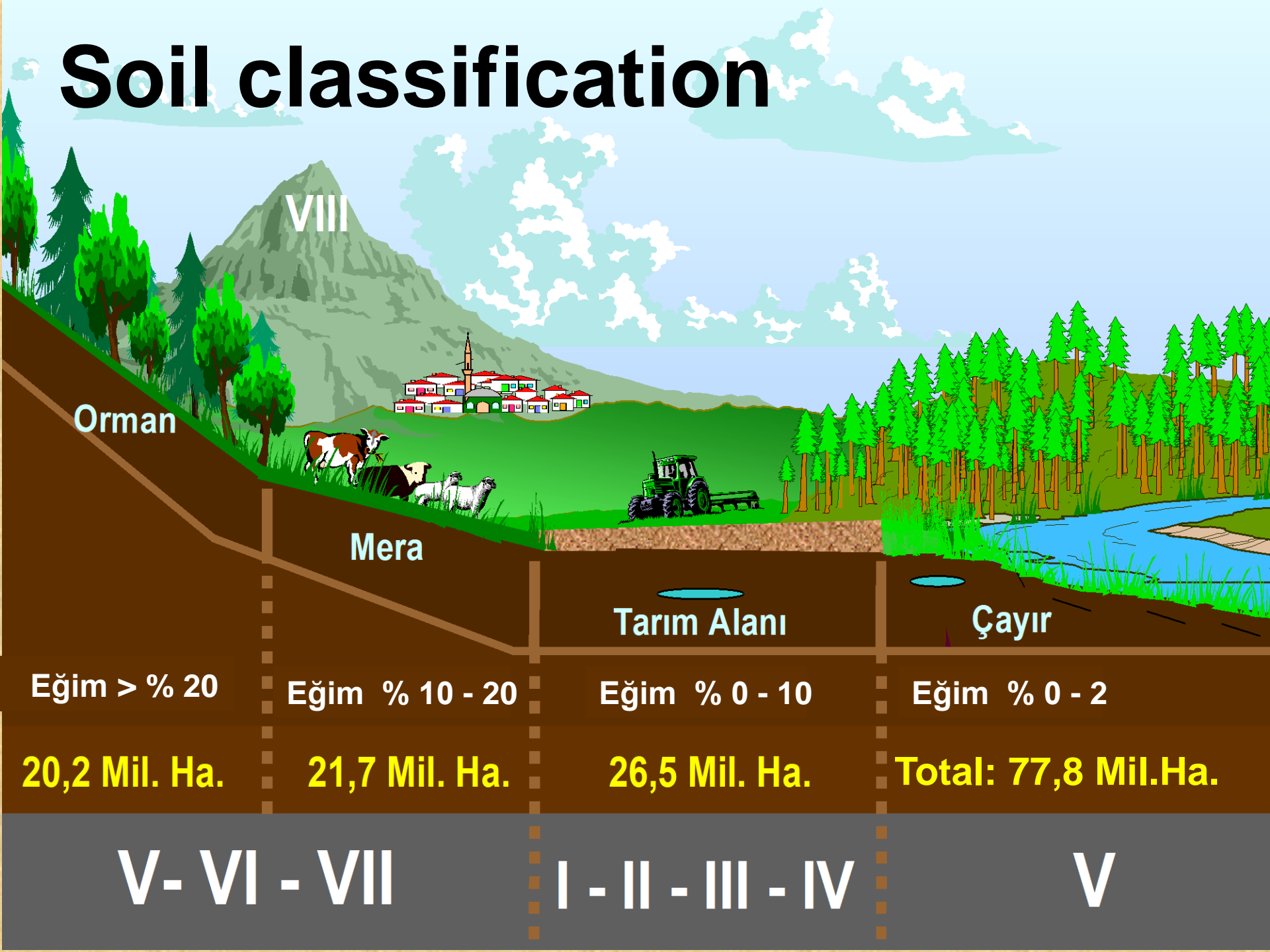
10.0 kV

10µm

TR16

#0026

Soil classification



VIII

Orman

Mera

Tarım Alanı

Çayır

Eğim > % 20

Eğim % 10 - 20

Eğim % 0 - 10

Eğim % 0 - 2

20,2 Mil. Ha.

21,7 Mil. Ha.

26,5 Mil. Ha.

Total: 77,8 Mil. Ha.

V - VI - VII

I - II - III - IV

V