

# MARINE AND OCEAN CHEMISTRY

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# PLAN – CONTENT – REFERENCES

1. Introduction
2. The water in seawater
3. Salinity, chlorinity, conductivity, and density
4. Major constituents of seawater
5. Simple gases
6. Salts in solution
7. Carbon dioxide
8. Nutrients
9. Trace metals and other minor elements
10. Chemical extraction of useful substances from the sea

## **References:**

1. An Introduction to the Chemistry of the Sea, Michael E. Q. Pilson
2. Marine Chemistry & Geochemistry, John H. Steele et al.
3. Chemistry in the Marine Environment, R. E. Hester and R. M. Harrison
4. Marine Chemistry, P. J. Wangersky

# SIMPLE GASES

1. Simple gas laws
2. Solubility in water
3. Sources and sinks within the ocean
4. Atmospheric exchange by diffusion

### Concentrations of conservative gases in dry air, the solubility of each pure gas in pure water

Gas	Symbol	Mole fraction in air	Solubility, $\mu\text{mol kg}^{-1} \text{ bar}^{-1}$	Concentration $\text{mol kg}^{-1}$
Nitrogen	N <sub>2</sub>	0.7808	695.0	$537.4 \times 10^{-6}$
Oxygen	O <sub>2</sub>	0.2095	1372.2	$284.7 \times 10^{-6}$
Argon	Ar	0.00934	1508	$13.95 \times 10^{-6}$
Neon	Ne	$18.2 \times 10^{-6}$	462.0	$8.328 \times 10^{-6}$
Helium	He	$5.25 \times 10^{-6}$	385.4	$2.00 \times 10^{-6}$
Krypton	Kr	$1.14 \times 10^{-6}$	2764	$3.12 \times 10^{-6}$
Xenon	Xe	$0.087 \times 10^{-6}$	4996	$0.434 \times 10^{-6}$

- The recommended SI unit of pressure is the Pascal, which is one Newton per square meter.
- The use of the bar ( $= 10^5$  Pa) is in common use, and has the advantage that 1 bar is close to the standard atmospheric pressure.

# SIMPLE GAS LAWS

For air:

$$p_T = pp(N_2) + pp(O_2) + pp(Ar) + pp(H_2O) + \dots$$

$p_T$  = The total pressure

$pp$  = The partial pressure



- Ideal gas law:

$$p V = n R T$$

- At STP (standard temperature and pressure: 0 °C or 273.15 K, and 1 atmosphere pressure = 101 325 Pa) 1 mole of an ideal gas occupies a volume of 22.413 L.
- All real gases deviate slightly from the “ideal” behavior.
- For precise work, it is sometimes necessary to account for these deviations by using one of several possible equations of state with additional empirically measured terms, such as the so-called “virial equation”:



$$P V_m = R T \left( 1 + \frac{B_T}{V_m} + \dots \right)$$

$V_m$  = The molar volume

$B_T$  = The second of several virial coefficients (a function of temperature and is different for each gas)

- For accurate calculation of the situations that might be found under high pressures in the ocean, the virial equation or some other formulation is necessary.

# SOLUBILITY IN WATER

- The amount of a gas dissolved in a quantity of water, when the gas and water phases are in equilibrium, is directly proportional to the pressure of the gas.
- Henry's law:

$$[G] = H_G \times pp(G)$$

$[G]$  = The concentration of gas  $G$  in solution

$pp(G)$  = The partial pressure of gas  $G$

$H_G$  = The Henry's law constant for gas  $G$  (a function of  $T$  and  $S$ )

- Since the major conservative gases in the atmosphere are present in very nearly constant proportions, it has been the usual oceanographic practice not to calculate from the Henry's law constant for each gas, or to give solubilities as concentrations in equilibrium with the pure gas.
- Instead, their solubilities are provided by tabulating their concentrations in seawater in equilibrium with the atmosphere, at 1 atmosphere total pressure, **under specified conditions of temperature and salinity.**

## Concentrations of three major atmospheric gases in water and seawater

Gas	Air, %	Pure Water		Seawater, $S = 35\text{‰}$	
		0 °C	30 °C	0 °C	30 °C
Concentrations ( $\mu\text{mol kg}^{-1}$ )					
N <sub>2</sub>	-	830.5	457.9	622.0	362.5
O <sub>2</sub>	-	457.0	237.3	347.9	190.7
Ar	-	22.30	11.65	17.02	9.38
Composition of total gas (mol%)					
N <sub>2</sub>	78.08	63.4	64.8	63.02	64.4
O <sub>2</sub>	20.95	34.9	33.6	35.3	33.9
Ar	0.934	1.70	1.65	1.72	1.67

# SOURCES AND SINKS WITHIN THE OCEAN

- In addition to gain or loss by exchange with the atmosphere, there are two possibilities.
  1. Gases may be produced or consumed in situ.
  2. They may exchange with the sediments.

# ATMOSPHERIC EXCHANGE BY DIFFUSION

- Both the air and the water are always in motion to some extent. Turbulent eddies commonly must keep both air and water mixed and chemically homogeneous at some distance above and below the interface.
- However, such eddies cannot penetrate exactly to the surface, and indeed, at some distance, perhaps very close to the interface, eddy motions must nearly vanish.
- A common way to imagine the situation is to consider that right at the interface there are **unstirred laminar boundary layers** that do not readily mix with the air<sub>4</sub> above or the water below.

- For non-reactive gases, such as  $O_2$ ,  $N_2$ , and the noble gases, it can be shown that they diffuse rapidly enough through the unstirred laminar boundary layer in the air that this is not a significant barrier. However, there is certainly a barrier imposed by the layer of unstirred water.
- For very reactive gases, such as  $SO_2$  or  $NH_3$ , the reaction rate at the water surface is so great that diffusion in the boundary layer of the air limits transport rates so in such a situation this must also be considered.